A SEDIMENTOLOGICAL STUDY OF THE CAP ENRAGE CONGLOMERATE, QUEBEC

TRANSPORT OF CONGLOMERATE INTO DEEP WATER: A STUDY OF THE CAMBRO-ORDOVICIAN CAP ENRAGE CONGLOMERATE AT ST. SIMON DE RIMOUSKI, QUÉBEC

By

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SCOPE AND CONTENTS: The Cambro-Ordovician sequence at St. Simon, Québec, was divided informally into ten horizons by Mathey (1970). The most prominent lithologies within the horizons are pelites, feldspathic sandstones and petromict conglomerates. One horizon 50 metres thick of feldspathic sandstones and conglomerates, described in this study, consists of three large fining upward sequences. The fining upward sequences are defined by the occurrence of five facies. These facies are: poorly sorted coarse conglomerates, well sorted coarse conglomerates, medium conglomerates with scattered pebbles and boulders, fine conglomerate with scattered pebbles and boulders and coarse sandstones. The base of each sequence is characterized by the occurrence of fine conglomerates and coarse sandstones. Rarely do the coarse conglomerates grade into thick developments of medium and fine conglomerates. The fine conglomerates grade in places into coarse sandstones, although generally the coarse sandstones

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have sharp bases. The conglomerates display sharp bases, normal and inverse grading, grain imbrication and orientation. The long axes of the grains, which define the orientation are parallel to, and not transverse to the flow direction suggested by the imbrication. To produce these features it is suggested that turbulence and dispersive pressures were operative within the flow. If the pebbles had moved as bed load material, they would have come to rest with their long axes transverse and not parallel to the flow direction suggested by the imbrication. The term "fluxoturbidite" (Dzulynski et al., 1959) has been applied to some coarse grained deposits in geosynclinal sequences. The characteristic features of "fluxoturbidites" are their unusually coarse grain size, thick irregular bedding with associated slump structures and poorly developed grading. The differences between the conglomerates described in this study and the typical "fluxoturbidite" preclude the use of this term to describe the deposits described by the author. It is suggested that the conglomerates were deposited upon a submarine fan complex by currents which flowed parallel to the present tectonic axes.

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CHAPTER 1

INTRODUCTION

Problems of Conglomerates

Conglomerates are common in the Cambro-Ordovician sequences which crop out from Vermont to Newfoundland. Conglomerates have also been reported and well described from many orogenic belts, particularly in Poland (Unrug, 1963), S. America (Scott, 1966) and Wales (Kelling and Wollards, 1969). In the Appalachian literature, many authors have mentioned the occurrence of conglomerates but few have provided detailed descriptions of the conglomerates with palaeocurrent directions.

Within the past twenty years there has been an advance in the level of understanding of deep water clastic sedimentation. Many geologists have been concerned with the mechanisms of transportation and deposition of sand into deep water environments but only a few have attempted studies of conglomerates. This is a result of several factors. Conglomerates are relatively uncommon in the geological record and they are intractable experimentally. Also, geologists have been unable to decide which attributes of conglomerates are the most important as indicators of transport mechanisms. Obviously the sedimentary structures associated with conglomerate deposits can provide some information about the origin of these deposits. Grain orientation and imbrication, and cross-stratification can provide palaeocurrent directions and possibly information on the mode of deposition of the rocks.

Several mechanisms have been proposed for the emplacement of coarse grained deposits in deep water environments. Some authors have stressed the importance of fluid turbulence, whereas others have stressed slump-slide mechanisms. The formulation of a model of deposition for any rock type must take into account the sedimentological characteristics of the rock type. In this study, emphasis has been placed on the description of the sedimentological characteristics of a conglomerate deposit, in the hope that these characteristics will provide some information on the importance of either fluid turbulence or slump-slide mechanisms during the deposition of the rocks.

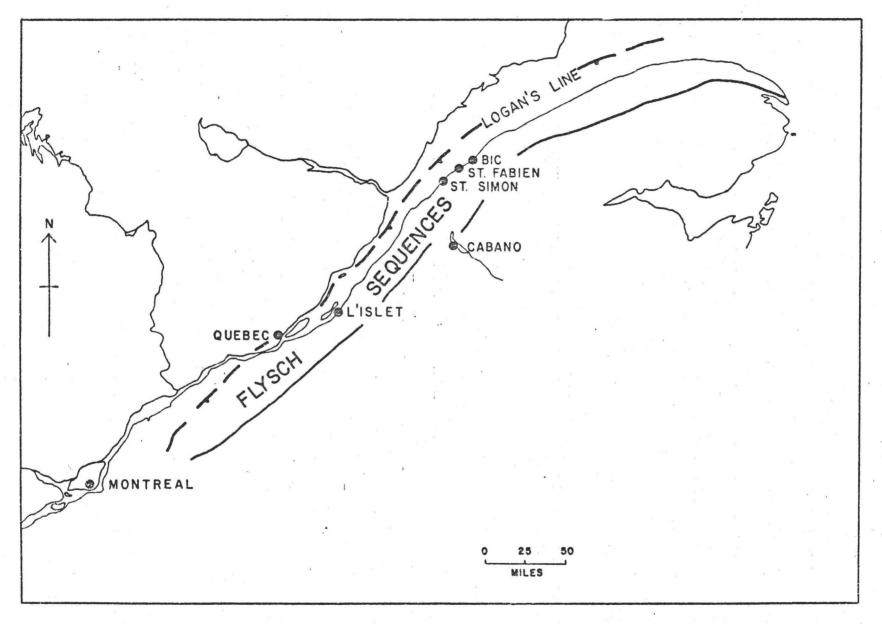
The conglomerate outcrop at St. Simon (Fig. 1) was chosen because a previous study (Mathey, 1970) had established the stratigraphical succession and nomenclature. Also, the exposure was such that lateral and vertical changes in the conglomerates could be easily seen.

Previous Studies of Cambro-Ordovician Conglomerates in Québec

Previous papers on the Cambro-Ordovician rocks in the Appalachians (Bailey and McInnes, 1893; Laverdiere and Morin, 1941), provided some descriptions of the conglomerates of the Québec Group which crop out along the south shore of the St. Lawrence. Osborne (1956) described the Québec Group in the Québec City area. He thought the conglomerates were originally deposited in shallow water, but were later resedimented into a deep water environment. He postulated that the site of origin of the conglomerates was either on the margin of the geosyncline or on unstable ridges within the geosyncline.

Hubert (1965) described some conglomerates in the Québec Group of the L'Islet-Kamouraska area (Fig. 1). The geometry of the conglomerate

Figure 1: Outline of the regional geology in the southern part of Québec. (Modified from Hubert et al., 1970.).



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outcrops (in the St. Roch and St. Damase Formations) and the distribution of maximum and mean grain sizes within them, suggested that the conglomerates were derived from the northwest. Palaeocurrent measurements taken from cross-bedding in the conglomerates and associated quartz bearing arkosic conglomerates indicate that the currents flowed from the northwest. Palaeocurrent measurements taken from cross-bedding and flute casts in the arkosic Armagh Formation at a lower level in the Québec Group, indicate that the currents flowed from the southeast and occasionally from the south and east.

Lesperance and Greiner (1969), recognized two conglomerate types in the Squatec-Cabano area (Fig. 1). One conglomerate type was thin bedded and interbedded with slates and siltstones, whereas the other was thick bedded and not associated with slates and siltstones. The thin bedded conglomerates possessed clasts with low sphericity and poor rounding. The thick bedded conglomerates possessed a well sorted matrix. Because of these differences they concluded that the thin bedded conglomerates were emplaced by a slide mechanism whereas the thick bedded ones were emplaced by turbidity currents.

Hubert et al. (1970) described two conglomerate deposits in the Québec Group. At L'Islet Wharf (Fig. 1), they described a channellized conglomerate-sandstone sequence which was deposited contemporaneously with a limestone-shale assemblage. Flute marks and channels in the conglomerate-sandstone assemblage indicate that the currents flowed from the north. In the Bic-St. Fabien area (Fig. 1), a conglomerate-sandstone assemblage (the Cap Enragé Formation) is underlain by at least 600 m of claystone and siltstone and is overlain by 1600 m of interbedded pelites

and graded sandstones. The overall stratigraphic position suggests the conglomerate-sandstone sequence was deposited in deep water. Palaeocurrent directions indicate flow from the north. This study was also done on the Cap Enragé conglomerates, but evidence will be presented which suggests that palaeocurrents flowing towards the southwest were also operative when the conglomerate-sandstone assemblage was deposited.

Mathey (1970) described the geological succession of the Cap Enragé Formation at St. Simon, about twenty miles S.E. of Bic (Fig. 1). He considered the lower part of the St. Simon section (essentially mudstones, feldspathic sandstones and conglomerates) to be laterally equivalent to the St. Fabien sequence described by Hubert et al. (1970) (Table 1, p. 6). Mathey's palaeocurrent information for the Cap Enragé Formation in the St. Simon district was obtained from cross-stratification, channels and flutes in the associated feldspathic sandstones. Few readings were obtained from the conglomerate assemblages themselves. He recorded twenty-two measurements in which the direction and sense of the currents were known and nineteen in which only the directions were known. Mathey's palaeocurrent information for the entire section is shown in Table 2, p. 7. The vector mean (for the succession) from the twenty-two measurements (direction and sense known) was 172 degrees, that from the nineteen measurements (direction known) was 144 degrees. This is contrary to the palaeocurrent information on the Horizon 3 conglomerates presented here, which indicates flow from the northeast, parallel to, rather than perpendicular to the present tectonic trends.

It should be stressed that throughout this thesis Mathey's stratigraphic nomenclature will be followed. A comparison of all the

TABLE 1. COMPARISON OF THE STRATIGRAPHICAL NOMENCLATURE

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Age	Lithologies	Stratigraphy At (Hubert et al.,		Stratigraphy At (Mathey, 1970)	t St. Simon	Stratigraphy At Bi (Lajoie In Press)
	Interbedded Pelites				8	Ladriére
	and Graded Sandstones	Ladriére	1	Niveaux	7	Formation
Lower		Formation			6	
Ordovician	7	5			5	Cap Enragé
	Quartzo-feldspathic	Cap Enragé 4		Niveau	4	Formation 3
	Sandstones	Formation				
laximum	Petromict Conglom-	Cap Enragé 3	a R	Niveau	3	Cap Enragé 2
ge	erates and Quartzo-	Formation 2				Formation
liddle	feldspathic Sandstones		,			
Cambrian	Quartzo-feldspathic	Cap Enragé 1		Niveau	2	Cap Enragé 1
	Sandstones	Formation				Formation
	Claystones and	Original		Niveau	1	Original
	Siltstones	Formation		*		Formation

TABLE 2. PALAEOCURRENT DATA FOR ST. SIMON

(Tabulated from Mathey, 1970)

Horizon	Direction an	d Sense Kno	wn	Directi	on Known
	Cross-Stratification	Channels	Flute Cast	Channels	"Flutes
8	160,150			den den den ser alle so boorden die er boorden	an a
7	100			×.	170
6	160,160,	185			
	160,140				
5				130,120,	
				010,010,	
				005	
4	230,260,				
	220				
3	200,185	180	•	160,150,	
		- 1		150,125,	
				020,015	
2		235,250,	140	170,160,	
		220,140		010,010,	
				000	
1	120,120,			130	140
	115				

schemes is given in Table 1.

Mathey (1970) recognized the following rock fragments in the conglomerates and sandstones at St. Simon; micritic, oolitic, sandy and

glauconitic limestones, quartzo-feldspathic sandstone, diabase and gneiss. The sandstone beds contained (in decreasing order of abundance) quartz, potassium feldspar, plagioclase feldspars, micas, glauconite, zircon, tourmaline and garnet. The igneous and metamorphic rock fragments and the heavy minerals are similar to the assemblages present in the Grenville Province of the Canadian Shield. The sedimentary rock fragments are similar to the Cambrian rocks which crop out south of the St. Lawrence River. Mathey concludes that the conglomerate-sandstone assemblages at St. Simon were formed by currents which resedimented material derived from the Canadian Shield and a carbonate platform in an area to the north of St. Simon. Mathey did not decide on a mechanism for the deposition of the coarse grained deposits but he suggests that they form a part of a submarine fan.

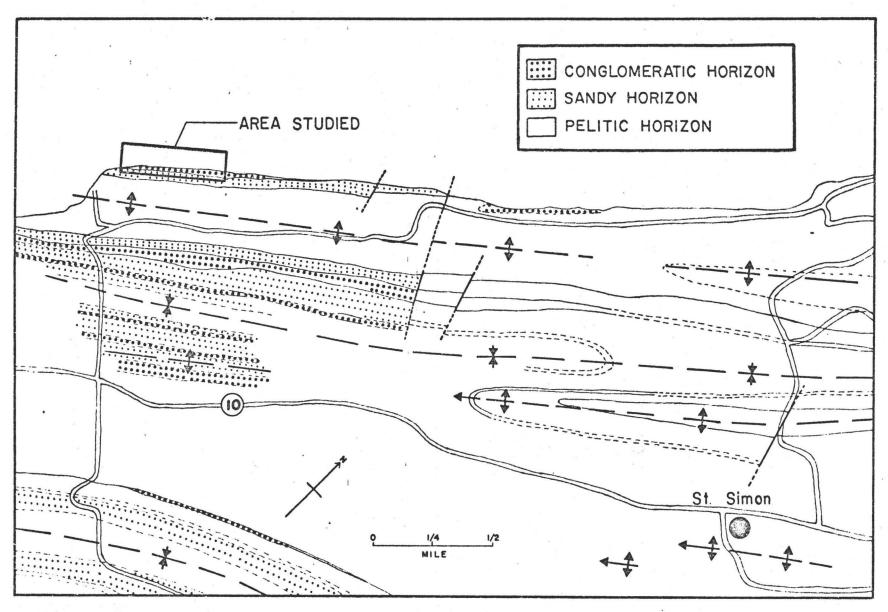
A subsidiary purpose of the present study is to fit the palaeocurrent information from St. Simon into a larger Appalachian picture of the environment of deposition. Models have been proposed by Dewey and Bird (1970) and Hubert et al. (1970). For Lower Palaeozoic times, Dewey and Bird (1970) envisaged a wide continental shelf bordered by a low craton in the west. Hubert et al. (1970) proposed a slightly different model of sedimentation for the rocks at L'Islet and St. Fabien. For Cambro-Ordovician times they proposed a narrow continental shelf, cut by submarine canyons, and bordered by a region of fairly high relief. The mountainous region supplied quartzo-feldspathic material which was rounded on the narrow continental shelf. Periodically this material was carried down the submarine canyons and deposited on submarine fan complexes. The palaeocurrent information presented here suggests that

parts of the Cambro-Ordovician fan complexes were constructed by currents which flowed parallel to the present tectonic axes.

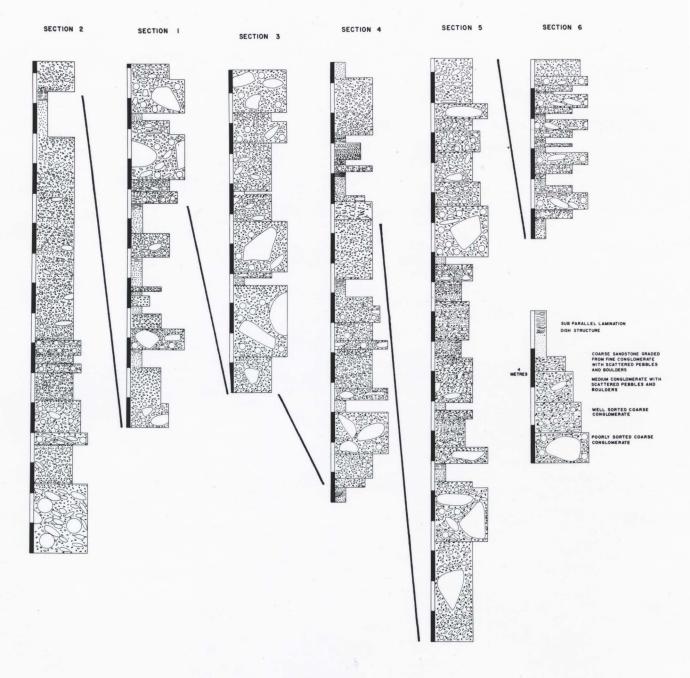
The Area of Study

This study deals with the sedimentological characteristics of the Cap Enrage Formation (Quebec Group) at St. Simon (Figs. 1 and 2). The Québec Group is part of the flysch sequence which crops out in the Gaspé Peninsula. In this area the flysch sequence is contained in a series of thrust slices. The age of the rocks within the thrust slices ranges from Lower Cambrian to Middle Ordovician. The sequences of rocks in the thrust slices are successively older from Logan's Line southward, the Middle Ordovician lying structurally below the Lower Ordovician, which is in turn covered by the Cambrian rocks (Hubert et al., 1970). The sequence in the St. Simon area is 600-700 m thick, and consists of an alternating series of conglomerate and feldspathic sandstone assemblages, with quartz sandstone assemblages and mudstone-siltstone assemblages. The sequence was subdivided by Mathey (1970) into ten informal stratigraphic horizons (Table 1). Horizon 3, described here, consists of conglomerates and feldspathic sandstones, with no siltstones and mudstones. It can be traced laterally from Trois Pistoles to Rimouski (40 miles). Perpendicular to strike the formation can be traced six miles south of the St. Lawrence River (Hubert et al., 1970). At Bic (the type area) the conglomerates of Mathey's Horizon 3 of the Cap Enrage Formation are about 18 m thick, but at St. Simon Horizon 3 is about 50 m thick.

The area is folded into tight parallel folds, whose axes trend S.W.-N.E. The outcrop described here is on the northern limb of one such fold, which has negligible plunge (Fig. 2). Along this limb, the beds Figure 2: General geology of the St. Simon area. (Modified from Mathey, 1970.).



generally dip to the northwest at a high angle. Some beds are overturned and dip to the southeast. There are numerous small faults trending S.E.-N.W., whose throws vary from several centimetres to a few metres (Map 1). Figure 3: Measured sections at six localities on the outcrop at St. Simon. The correlations are based upon facies similarity and the "walking out" of beds on the outcrop.





CHAPTER 2

FIELDWORK

Mapping

Because a detailed study of the area was required, it was decided to map the outcrop at a scale of 1:100. An arbitrary zero bench mark was set up at the western end of the outcrop and distances were measured in metres along strike from this point. At intervals of 8-15 metres along the outcrop, substations were mapped by compass and tape and marked with white paint. These substations were used to locate positions on the outcrop.

At various distances along strike, stratigraphic sections through the sequence were measured (Figure 3). The position of these sections, and of photographs and samples, are shown on the map (Map 2).

During mapping, it became apparent that there were several facies exposed. The facies range from coarse conglomerates to fine conglomerates and sandstones (Chapter 3). Combinations of the facies define three fining upward sequences (compare Maps 1 and 2). Coarse conglomerates are prominent at the base of each fining upward sequence, whereas at the top of each sequence fine conglomerates and sandstones are prominent. The fieldwork carried out was aimed at determining the character and extent of the facies within the fining upward sequences.

Grain Orientations

After a reconnaissance of the outcrop, made during the first

field mapping season, it was apparent that many of the beds contained grains with a preferred a-axis orientation. On bedding surfaces the preferred a-axis orientation is called "grain orientation" and in sections perpendicular to bedding it is called "imbrication". Reconnaissance also showed that grain orientation and imbrication data could be obtained from photographs. A large number of surfaces in the outcrop could not be used for orientation studies for several reasons. Some surfaces were cleaved, others were oblique to bedding and other surfaces lacked contrast. Those surfaces which were covered with slime were cleaned with a wire brush before they were photographed. Where it was decided that a surface was photographable, the directions of dip and strike and a suitable scale were drawn on the surface. The location of the surface, the direction and amount of dip, the way up direction and the type of surface were recorded. Two faces perpendicular to bedding, one parallel to strike and one transverse to strike, were examined in order to estimate the 3-D imbrication as close as possible to the location of each orientation photograph. In some cases, this was not possible. The surfaces were photographed at a distance of approximately one metre.

20 x 25 cm prints were made and covered with tracing paper. Equispaced lines, which did not intersect the same grain twice, were drawn at right angles to the direction of strike. The print and the affixed tracing paper were mounted on a light table. The long axes of grains which fell on the equispaced lines were drawn according to the least projection elongation method. The long axes were drawn only for grains which were not broken and which possessed sharp boundaries. (Elongate grains are those which possess an apparent axial ratio a/b

greater than 1.5.). After one hundred long axes had been drawn by the author, the tracing paper was removed from the print and the angles of the long axes were measured with reference to strike by student assistants. A vector mean and vector strength were computed for each sample.

Imbrication

Reconnaissance showed that only those faces roughly parallel to the present strike exhibited imbrication. No preferred pebble fabric could be demonstrated on faces perpendicular to the present strike. To record imbrication information, a tape measure was placed at the top of the bed in such a way that the centimetre values on the tape decreased eastwards. The tape was parallel to regional bedding. The surfaces were photographed at a distance of approximately one metre and the location of each photograph was recorded. The long axes of the grains were drawn using the same method as outlined for the grain orientations.

Two sets of data were recorded. The first set contained one hundred readings obtained from all elongate grains regardless of size. (An elongate grain is one which possesses an apparent axial ratio a/c greater than 1.5.). The second set was obtained from either fifty or twenty-five elongate grains longer than three centimetres. The number of grains measured in the second set of data was controlled by the number of grains on the print. If there were less than fifty grains on the print, only twenty-five were measured. The vector mean and vector strength were computed for each outcrop.

Roundness

After examination of the outcrop, it was decided that grain roundness measurements could be obtained by using the method of Dobkins and Folk (1970). On a suitable basal surface the apparent roundness of a grain was measured by comparing the diameter of curvature of the sharpest corner visible with the diameter of the largest inscribed circle to fit the grain. One hundred apparent roundness measurements were calculated at each outcrop (Appendix 1). In this work, all the roundness measurements quoted are apparent roundness measurements.

Elongation

The apparent elongation of a grain is computed as:

apparent length of a-axis apparent length of c-axis.

A grid system for sampling was set up at several localities and the apparent elongation of at least one hundred grains was calculated (Appendix 2).

Measurements of the Ten Largest Clasts

In order to see if there was an appreciable change of maximum clast size up through the section, and if there was any relationship between clast size and lithology, it was decided to measure the diameter of the ten largest limestone and ten largest sandstone clasts present at various horizons. The mean size of the ten largest clasts was regarded as an approximation to the coarsest material that could be transported by the currents. For convenience, the "average diameter of the ten largest clasts" will be symbolized as D/10.

Composition

The compositions of the conglomerates were estimated using a grid system. The grid consisted of parallel equispaced lines, of equal length, which were drawn on the outcrop. The lithology and the length of intersection of each grain (with an apparent a-axis longer than 1 mm) along the lines of the grid system were recorded. Those grains with an apparent a-axis less than 1 mm were recorded as matrix. The total lengths of the grid systems differed from facies to facies. In coarse conglomerates the lengths of the grid systems were about 5 m, in the medium conglomerates the lengths were about 3 m.

CHAPTER 3

Facies Definition and Description

During field mapping it was noticed that there was a wide range of clast size in the conglomerates. The clast size (apparent length of the a-axis) seemed to be the most characteristic property of the conglomerates and for this reason it was decided to subdivide the sequence into four lithologies;

> coarse conglomerate, a-axis > 4 cm medium conglomerate, a-axis 1.5-4 cm fine conglomerate, a-axis 0.5-1.5 cm coarse sandstone, a-axis < 0.5 cm.

The range of clast size, degree of sorting, and the amount of coarse sandstone within a certain stratigraphic thickness were sufficient to define a "facies", that is, a restricted part of the stratigraphic succession which is different in aspect from adjacent parts of the succession. The facies are:

> Coarse Conglomerate, Poorly Sorted Coarse Conglomerate, Well Sorted Medium Conglomerate with Scattered Pebbles and Boulders¹

¹In this work pebbles are clasts which have an apparent a-axis 4-25 cm long and boulders have an apparent a-axis > 25 cm long.

Fine Conglomerate with Scattered Pebbles and Boulders Coarse Sandstones.

The poorly sorted conglomerates have more, and larger boulders than the well sorted conglomerates (compare Figs. 8 and 24). The types of occurrence of the facies are shown in Table 3, p. 20.

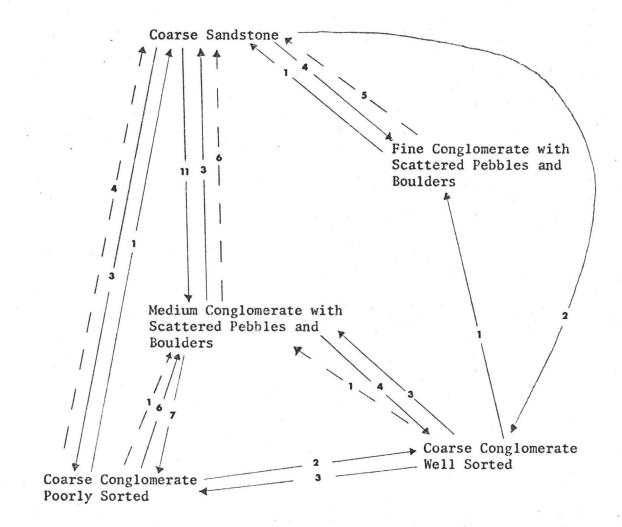
The coarse and medium conglomerate beds commonly have thin (< 10 cm) layers of finer grain size at the top. These thin layers were not mapped individually and are not included in Table 3, although they are described under the appropriate facies headings. Lateral variations in the facies were not as prominent as the vertical ones. The concentration of coarse conglomerate clasts varied laterally in some units. In such cases the prominent lithology defined the facies for basic mapping and descriptive purposes.

In this work, within-facies layering shown up by changes in the proportion of different grain sizes is defined as stratification. The stratifications in the medium conglomerate are generally thicker than those in the fine conglomerates.

It was noticed that combinations of the facies defined three overall fining upward sequences (compare Maps 1 and 2). Several generalizations can be drawn about these sequences. Either or both of the coarse conglomerate facies occur near the base. The percentage of coarse conglomerate decreases upwards in the first and second sequence but near the top of the third, thin coarse conglomerates occur. The coarse conglomerates are usually succeeded by medium conglomerates. In the first and second sequences the percentage of medium and fine conglomerate increases upwards. The top of each sequence is marked by a thick

of The Facies	Basal Inverse Grading, Main Part Of The Bed Structureless	Basal Inverse Grading, Main Part Of The Bed Normally Graded	Normal And Inverse Grading At The Base, Main Part Of The Bed Normally Graded	Whole Bed Normally Graded	Whole Bed Stratified	Bed Partly Stratified	No Structure At All
· · ·			Normally Graded				
Coarse Conglomerate, Poorly Sorted					•		
Individual Beds Of Coarse Conglomerate	4	5		3			
Coarse Conglomerate Grading To Medium Conglomerate With Scattered Pebbles And Boulders			1				
Coarse Conglomerate Grading To Coarse Gandstone	4						
Coarse Conglomerate, Well Sorted		1					
Individual Beds Of Coarse Conglomerate	1	2	3	2	2		
Coarse Conglomerate Grading To Medium Conglomerate And Coarse Sandstone			1				
Medium Conglomerate Wit Scattered Pebbles And Boulders	h						
Scattered Pebbles And	h				16		13
Scattered Pebbles And Boulders Individual Beds Of	h				16		13
Scattered Pebbles And Boulders Individual Beds Of Medium Conglomerate Medium Conglomerate Grading To Fine Conglomerate And	h				16		
Scattered Pebbles And Boulders Individual Beds Of Medium Conglomerate Grading To Fine Conglomerate And Coarse Sandstone Medium Conglomerate Grading To Coarse	h						1
Scattered Pebbles And Boulders Individual Beds Of Medium Conglomerate Medium Conglomerate Grading To Fine Conglomerate And Coarse Sandstone Medium Conglomerate Grading To Coarse Sandstone Fine Conglomerate With Scattered Pebbles And	h						1
Scattered Pebbles And Boulders Individual Beds Of Medium Conglomerate Grading To Fine Conglomerate And Coarse Sandstone Medium Conglomerate Grading To Coarse Sandstone Fine Conglomerate With Scattered Pebbles And Boulders Individual Beds Of	h			4	4		1
Scattered Pebbles And Boulders Individual Beds Of Medium Conglomerate Medium Conglomerate Grading To Fine Conglomerate And Coarse Sandstone Medium Conglomerate Grading To Coarse Sandstone Fine Conglomerate With Scattered Pebbles And Boulders Individual Beds Of Fine Conglomerate Fine Conglomerate Fine Conglomerate Grading To Coarse	h			4	4		1

TABLE 3. TYPES OF OCCURRENCE OF THE FACIES



2→ Sharp Boundary Between Facies; Occurs Twice
 2→ Gradational Passage Between Facies; Occurs Twice

Figure 4: Facies relationship diagram for the entire section at St. Simon to show the type of boundary between the facies and the number of times the facies are in vertical contact with each other. (Modified from De Raaf et al., 1965.).

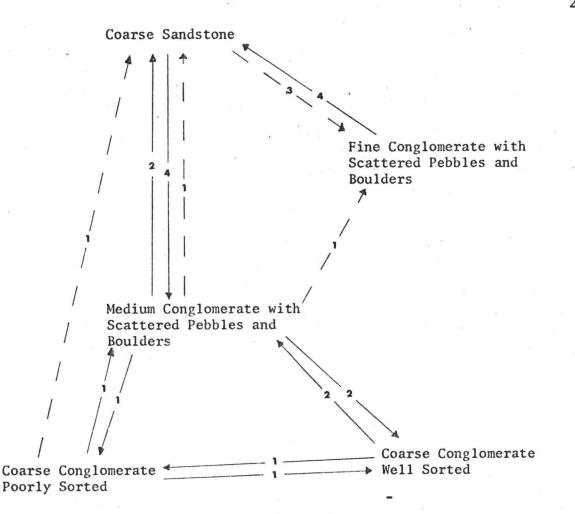
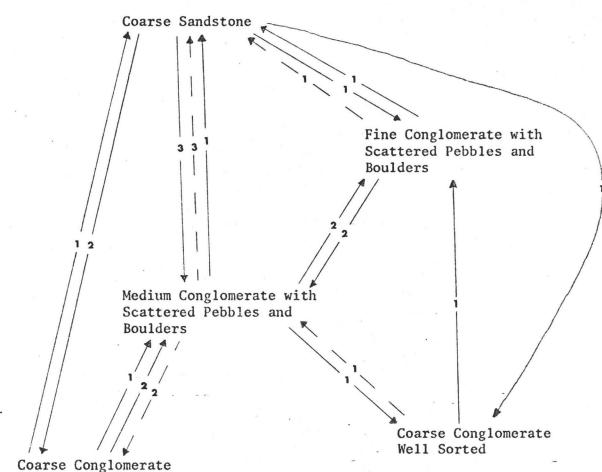


Figure 5: Facies relationship for the first fining upward sequence at St. Simon to show the type of boundary between the facies and the number of times the facies are in vertical contact with each other. (Modified from De Raaf et al., 1965.).



Coarse Conglomerate Poorly Sorted

— 2 → Sharp Boundary Between Facies; Occurs Twice
 — 2 → Gradational Passage Between Facies; Occurs Twice
 Figure 6: Facies relationship for the second fining upward sequence at St. Simon to show the type of boundary between the facies and the number of times the facies are in vertical contact with each other. (Modified from De Raaf et al., 1965.).

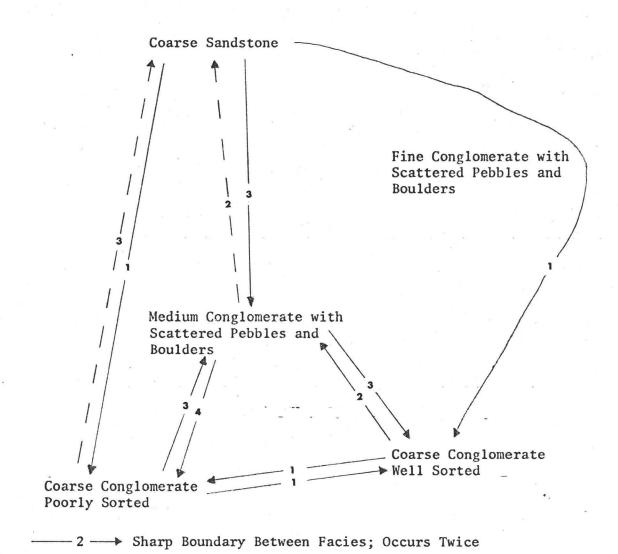


Figure 7: Facies relationship for the third fining upward sequence at St. Simon to show the type of boundary between the facies and the number of times the facies are in vertical contact with each other. (Modified from De Raaf et al., 1965.).

Gradational Passage Between Facies; Occurs Twice

-2 ----

sandstone, and in all the sequences the percentage of sandstone increases near the top. The thickest sandstones are in the first sequence, the thinnest and most persistent laterally are in the third. The facies relationships within the entire section and within each of the three fining upward sequences are shown in Figures 4-7, p. 21-24.

Coarse Conglomerate, Poorly Sorted

General Features

This facies comprises 25 percent of the entire section (Fig. 8). The types of occurrence of the facies are shown in Table 4, p. 26. The average bed thickness is about 80 cm, with beds ranging from 20 cm to 2 m (Map 1). Generally, the bases of the beds are flat and sharp, but 30 percent have channels or scours up to 1 m deep (Fig. 9). Individual beds are difficult to trace because of faulting. However, groups of beds (up to 4 m thick) can be correlated laterally for 400 m parallel to the inferred direction of flow. Within areas in which individual beds can be walked out, there is no evidence for "shingling", that is, the building forward of successive beds in partly-overlapping wedges. Twenty percent of the beds have load structures. The conglomerates load to a depth of 3-4 cm for several metres along the underlying sandstones (Fig. 10). The largest clast (250 cm) is in the first fining upward sequence (Bed 16B, Map 1) but the maximum D/10 (88 cm) is in the second fining upward sequence (Appendix 3, Bed 26, Fig. 11, Map 3). In the field about half the beds appear to have a bimodal grain size distribution, with modes of about 6 and 11 cm. A subordinate coarser modal class of about 20-30 cm was also noted. There is no correlation between clast size and lithology.

TABLE	4.	TYPES OF OCCURRENCE OF THE FACIES
		(Coarse Conglomerate, Poorly Sorted)

Types Of Occurrence Of The Facies	Basal Inverse Grading, Main Part Of The	Basal Inverse Grading, Main Part Of The	Normal And Inverse Grading At The Base,	Whole Bed Normally Graded
	Bed Structureless	Bed Normally Graded	Main Part Of The Bed Normally Graded	

4

4

Individual Beds Of Coarse Conglomerate

Coarse Conglomerate Grading To Medium Conglomerate With Scattered Pebbles And Boulders

Coarse Conglomerate Grading To Coarse Sandstone 5

3

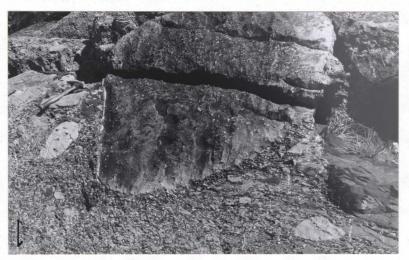


Figure 8: Outcrop of poorly sorted coarse conglomerate.



Figure 9: Channellized outcrop of poorly sorted coarse conglomerate.



Figure 10. Poorly sorted coarse conglomerate loaded in coarse sandstone.

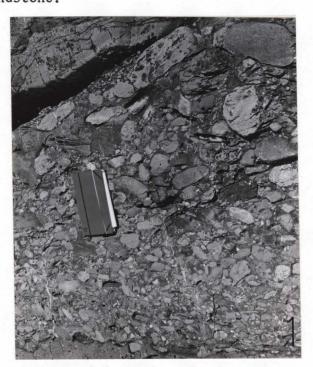


Figure 12: Inversely graded poorly sorted coarse conglomerate.

Figure 11: Distribution of D/10 measurements throughout the entire section.

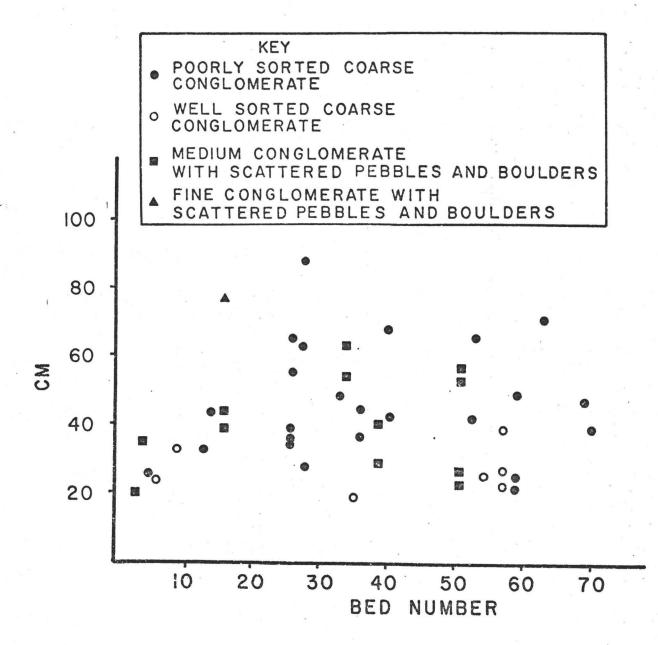


Figure 11

Internal Structures

The occurrence of the various internal structures is shown in Table 4, at the beginning of this facies description (p. 26). In those beds which display basal inverse grading, the grading is generally confined to within 15 cm of the base (Fig. 12), but in some beds it occurs up to 30 cm from the base. The inversely graded portion is medium conglomerate but pebbles and boulders occur. There is a gradual increase in the percentage of pebbles and boulders upwards, and the facies becomes a coarse conglomerate. Where the normal and inverse grading occur together, the inverse grading may extend laterally for 5 m before it is replaced by normal grading. The change from inverse to normal grading is shown up by an increase in the number of pebbles and boulders at the base of the bed. All the individual beds have thin developments of medium and fine conglomerates at their tops. The medium and fine conglomerates are not stratified.

The coarse portion of the facies is never stratified. A common feature of the coarse portion is the concentration of pebbles in linear or arcuate zones (Fig. 13). In the zones the pebbles are imbricate but do not overlap each other for any considerable length (less than 3 cm). The zones, which persist for one or two metres, dip east and west.

Composition

The composition of the facies, which is relatively constant, is summarized in Table 5, p. 31 (Figs. 14-17, Map 2, Appendix 4). The low percentage of matrix suggests the facies is grain supported. In one third of the beds the percentage of quartz increases near the top. The

Constituent	Range (%) Throughout The Section (Bed And Sequence Nos. In Parenthesis)	Range (%) Within The Sequences (Bed And Sequence Nos. In Parenthesis)
Sandstone	16(63,3) - 27(2,1)	16(63,3) - 21(65,3)
Quartz	1(65,3) - 9(36,2)	2(33,2) - 9(36,2)
Micrite	27(33,2) - 38(63,3)	28(65,3) - 38(63,3)
Matrix -	11(65,3) - 22(63,3)	12(33,2) - 15(36,2)

TABLE 5.SUMMARY OF THE COMPOSITION OF THE FACIES2(Coarse Conglomerate, Poorly Sorted)

 $^2\!A11$ those grains with apparent a-axes less than 1 mm were classed as matrix.

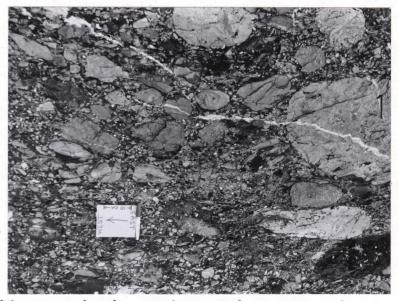


Figure 13: Pebble zones in the poorly sorted coarse conglomerate. The pebble zone is the arcuate feature extending from the top right-hand to the top left-hand corner of the figure.

Figure 14: Range of the percentage of micrite clasts contained within the facies.

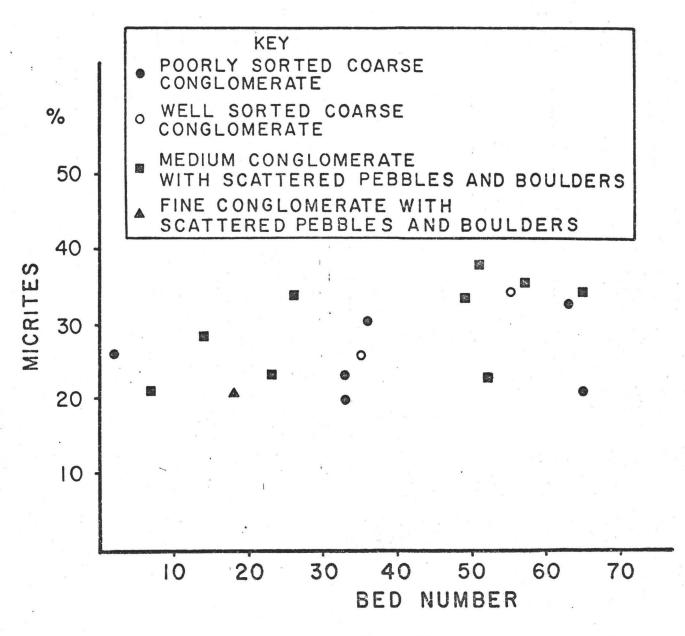


Figure 14

Figure 15: Range of the percentage of sandstone clasts contained within the facies.

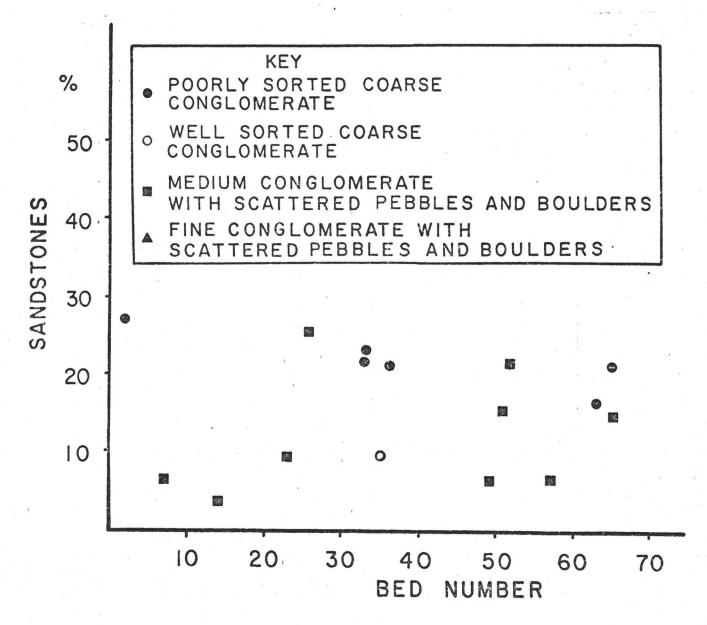


Figure 15

Figure 16: Range of the percentage of quartz clasts contained within the facies.

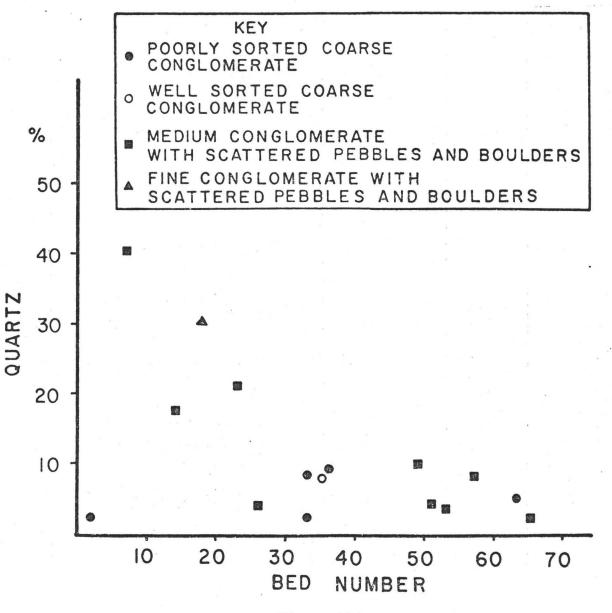


Figure 16

35

Figure 17: Range of the percentage of matrix contained within the facies.

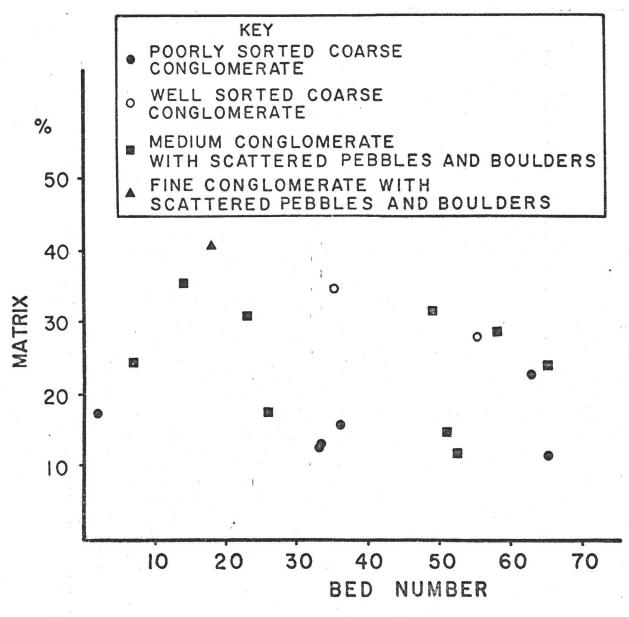


Figure 17

36

increase is a rapid one, occurring within 15 cm. The quartz grains near the base of the bed are coarser than those near the top. The lithologies of the coarse conglomerate clasts are; black and green meta-limestones, parallel and cross-laminated sandy limestones, grey micrites, pisolitic limestones, bioclastic limestones, calcite cemented medium grained cream sandstones with rounded quartz grains, calcite cemented medium grained cream sandstones and light grey fine grained sandstones. Some groups of beds are characterized by certain lithologies. Within the second fining upward sequence, one group (Beds 26-28, Map 1) characterized by cream sandstone clasts can be traced laterally for 400 m.

Roundness and Elongation

A roundness study of one hundred grains on the base of one bed at one locality gives an average apparent roundness of 0.40 (Dobkins and Folk, 1970) (Bed 1, Map 3, Appendix 1). Clast roundness ranges from well rounded to angular (Fig. 18).

Within one bed no consistent change of elongation occurs vertically. Also, no consistent change of the average elongation index (the average of the elongation measurements at one locality) occur throughout the section. The index ranges from 2.2 to 2.6 (Beds 63 and 36, Map 3, Appendix 2).

Imbrication

Measurements of the well developed easterly dipping imbrication do not reveal any apparent change of angle vertically or laterally within the section (Map 4, Figs. 19-20, Appendix 5). Generally, the imbricate grains are separated by smaller grains (apparent a-axis of 1 cm and less) and matrix (Fig. 21). Where imbricate grains are in contact, the contacts are usually less than 3 cm long. The measurements of imbrication are summarized in Table 6.

No. Of Grains	Range (In Degrees) Of Imbrication (Bed And Sequence Nos. In Parenthesis)
100	1(40,2) to 8(59,3)
50	6(40,2) to 14(26,2)

TABLE 6. SUMMARY OF THE IMBRICATION MEASUREMENTS (Coarse Conglomerate, Poorly Sorted)

The smaller grains possess a more random imbrication, filling the spaces between the larger clasts (Fig. 22). The direction of imbrication of the small clasts changes near very large clasts (Fig. 23). Within-bed angles of imbrication are laterally variable, ranging from 11 to 14 degrees and 6 to 8 degrees (Beds 26 and 40, Map 4). Towards the top of many beds, the imbrication dies out and flat lying grains increase in proportion. The flat lying grains are less than 3 cm long and are separated by coarse sand and matrix.

Three reversals of imbrication, that is, grains with their long axes dipping westward, occur in the second fining upward sequence (Bed 40, Map 4). The angles of imbrication of 2, 6 and 1 degrees are on the same bed but eight and 130 m apart respectively. Also on the same bed, imbrications (grains with long axes dipping eastward) of 1 degree (measured from one hundred grains), 6 and 8 degrees (measured from fifty and twenty-five grains respectively) are recorded. Another reversal of



Figure 18: Range of clast roundness within the poorly sorted coarse conglomerates.



Figure 21: Imbricated poorly sorted coarse conglomerate.

Figure 19: Imbrication measurements obtained from one hundred grains.

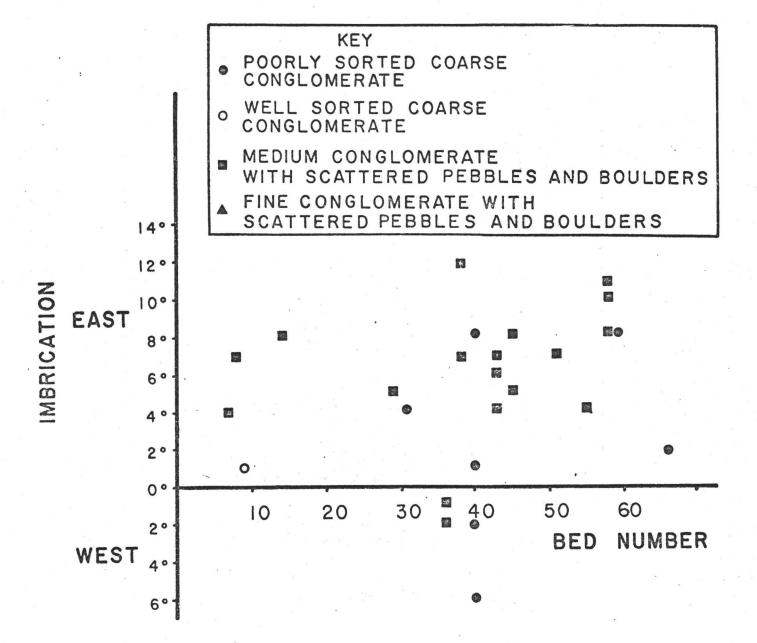




Figure 20: Imbrication measurements obtained from fifty and twenty-five grains.

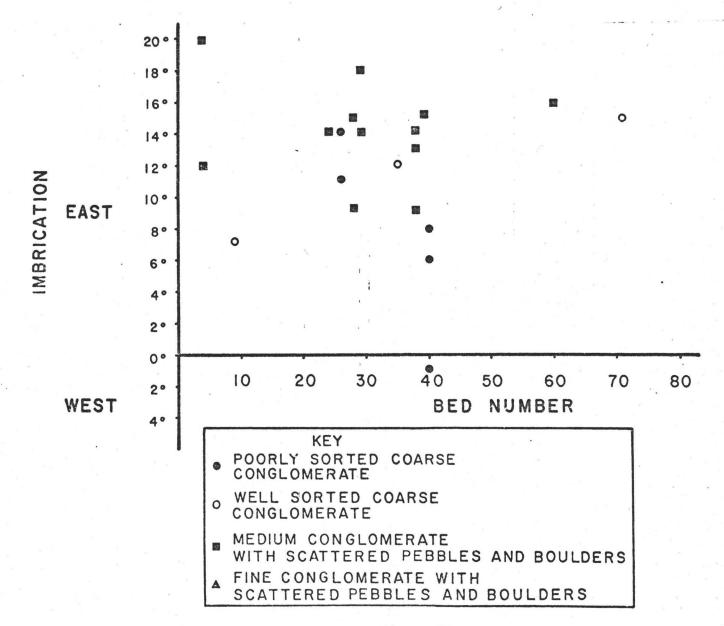


Figure 20

imbrication (long axis dipping westward) occurs in Bed 36 about 20 m east of those measured in Bed 40 (Map 4).

Coarse Conglomerate, Well Sorted

General Features

This facies comprises about 7 percent of the entire section (Fig. 24). The types of occurrence of the facies are shown in Table 7, p. 44. The average bed thickness is about 60 cm with beds ranging from 20 to 185 cm. Sixty percent of the beds have flat, sharp bases and 40 percent have channels or scours up to 60 cm deep. It is difficult to trace individual beds along strike because of faulting. The beds, which can be traced laterally for 40 m, do not appear to be shingled. No slump or load structures are visible in the facies.

The largest clast (76 cm) is in the third fining upward sequence (Bed 70) but the maximum D/10 is in the second fining upward sequence (39 cm) (Bed 57, Map 3, Fig. 11, Appendix 3). Measurements of D/10 range from 19 to 42 cm. There is no systematic change of the maximum clast size vertically through the section, nor is there any systematic change of D/10. Also there is no consistent change laterally (within any bed) of boulder size. In the field about half the beds appear to have a bimodal grain size distribution, with modes of about 6 and 11 cm. There is no relationship between maximum clast size and lithology.

Internal Structures

The occurrence of the various internal structures is shown in Table 7, at the beginning of this facies description (p. 43). In those beds which display basal inverse grading, the grading is generally TABLE 7.TYPES OF OCCURRENCE OF THE FACIES
(Coarse Conglomerate, Well Sorted)

Types Of Occurrence Of The Facies	Basal Inverse Grading, Main Part Of The Bed Structureless	Basal Inverse Grading, Main Part Of The Bed Normally Graded	Normal And Inverse Grading At The Base, Main Part Of The Bed Normally Graded	Whole Bed Normally Graded	No Structure At All
Individual Beds Of Coarse Conglomerate	1	2	3	2	2
Coarse Conglomerate Grading To Medium Conglomerate And Coarse Sandstone			1		



Figure 22: Small grain filling the space between larger clasts in the poorly sorted coarse conglomerate facies. The small clast occurs above the six inch mark on the tape.



Figure 23: Reversals of imbrication near a large clast in the poorly sorted coarse conglomerate facies. The long axes of the grains dip westward.

confined to within 20 cm of the base. The inversely graded portion is medium conglomerate but pebbles and boulders occur. There is a gradual increase in the percentage of pebbles and boulders upwards. In those beds which display normal and inverse grading, the grading is similar to that shown by the poorly sorted coarse conglomerates. With the exception of those beds which are not graded and the bed which grades into a medium conglomerate and coarse sandstone, the coarse conglomerates have thin developments of medium and fine conglomerates at their tops. Only in one bed (Bed 58) is the fine conglomerate stratified. The stratifications are subparallel and can be traced for 10 m.

The coarsest portion of the facies is never stratified. Linear pebble zones similar to those in the poorly sorted coarse conglomerates are common in this facies.

Composition

The composition of the facies is relatively constant throughout the section (Figs. 14-17, Appendix 4) and is summarized on Table 8.

TABLE	8.	SUMMARY	OF	THE	COMPOS	ITION	OF	THE	FACIES
		(Coarse	Con	ig1on	merate,	Well	Soi	rted)	

Constituent	Range (%) Throughout Section (Bed And Sequence Nos. In Parenthesis)				
Sandstone	1(55,3) - 9(35,2)				
Quartz	1(55,3) - 8(35,2)				
Micrite	26(35,2) - 34(55,3)				
Matrix	27(55,3) - 34(35,2)				

³All those grains with apparent a-axes less than 1 mm were classed as matrix.

The low percentage of matrix suggests the facies is grain supported.

Roundness and Elongation

Two roundness studies, each of one hundred grains, at different localities, give average roundness values of 0.42 (Bed 42) and 0.50 (Bed 23) (Dobkins and Folk, 1970) (Map 3, Appendix 1). These values are higher than that recorded for the poorly sorted conglomerate facies (0.39). Clast roundness ranges from well rounded to angular. The elongation measurements (obtained from two localities) do not show any significant changes vertically within a bed. The average elongation indexes (the averages of the elongation measurements at the two localities) are similar to those measured in the poorly sorted conglomerate, about 2.3 (Map 3, Appendix 2).

Imbrication

Measurements of the easterly dipping imbrication show an increase from the bottom to the top of the section. The increase is from 7 degrees in the first fining upward sequence to 12 degrees in the second to 15 degrees in the third (Beds 9, 35 and 71 respectively, Map 2) (Figs. 19-20, Appendix 5). Generally, the imbricate grains are separated by small grains and matrix (Fig. 25). The small grains (apparent a-axis of 1 cm or less) possess a more random orientation. Near the top of many beds the imbrication dies out and flat lying grains increase in proportion. The flat lying grains are generally less than 3 cm long and are separated by coarse sand and matrix. No reversals of imbrication were recorded in this facies.



Figure 24: Outcrop of well sorted coarse conglomerate.

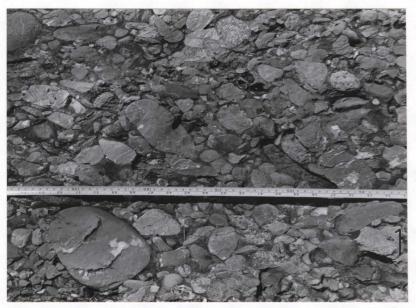


Figure 25: Imbricated well sorted coarse conglomerate. Numbers on the tape decrease eastwards.

Medium Conglomerate with Scattered Pebbles and Boulders

General Features

This facies comprises 45 percent of the sequence (Fig. 26). The types of occurrence of the facies are shown in Table 9.

Types Of Occurrences Of The Facies	Whole Bed Stratified	No Structure At All
Individual Beds Of		
Medium Conglomerate	17	13
Medium Conglomerate Grading To Fine		
Conglomerate And Coarse Sandstone		1
Medium Conglomerate Grading To Coarse		
Sandstone	4	1
Medium Conglomerate Graded From Coarse		
Conglomerates	2	

TABLE 9. TYPES OF OCCURRENCE OF THE FACIES (Medium Conglomerate)

The average bed thickness is 90 cm with beds ranging from 6 cm to 7 m. Twelve percent of the conglomerates have channelled or scoured bases. The channels are up to 2.5 m wide and 60 cm deep (Map 1) and commonly contain coarse conglomerate material. Twelve percent of the conglomerates have loaded into the sandstones which underlie them. The loads, are generally asymmetric, up to 170 cm long and 25 cm deep. The steeper eastern wall of the loads is sometimes overturned to enclose partially a pocket of conglomerate (Fig. 27). The load structures are developed near to or contain coarse clasts. Seventy percent of the conglomerates have sharp, flat or undulating bases. Individual beds are difficult to trace because of faulting. However, groups of beds up to 2 m thick can be correlated over 300 m. The faices does not display any slump structures. The largest clast (265 cm, Bed 34) is in the second fining upward sequence. The maximum D/10 (63 cm) is also in the second sequence (Bed 34, Fig. 11, Map 3, Appendix 3). The D/10 values do not show any consistent change vertically throughout the section nor laterally within a group of beds. In 20 percent of those conglomerates which are stratified, the coarsest grains are at the base. In 70 percent of the stratified conglomerates the pebbles and boulders are scattered haphazardly throughout. In the remaining 10 percent of the conglomerates the pebbles and boulders increase in size and percentage upwards. The pebbles appear to have a bimodal grain size distribution with modes of 6 cm and 9-11 cm. There is no relationship between clast size and lithology.

Internal Structures

Sixty percent of all the conglomerates are stratified throughout their entire thickness. The stratified conglomerates can be subdivided as follows:

25 percent have low angle stratification

75 percent have subparallel stratification.

The stratifications, which are shown up by the abundance of quartz granules, are up to 50 cm thick (Fig. 28).



Figure 26: Outcrop of Medium Conglomerate with Scattered Pebbles and Boulders.

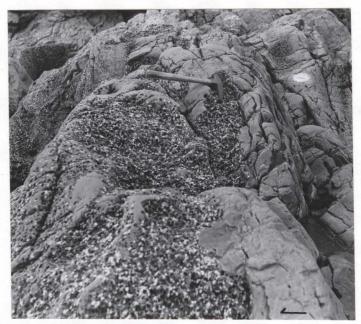


Figure 27: Load structures of Medium Conglomerate with Scattered Pebbles and Boulders in Coarse Sandstone. Twenty-five percent of the stratified conglomerates have low angle stratification with apparent dips up to 14° west and 14° east. The low angle stratification can occur as a single stratum or in sets up to 50 cm thick. Two types of sets are recognized. In one type the stratifications are straight, parallel and dip either east or west. In the other type the stratifications are concave upwards and dip east and west within the same unit (Fig. 29). Some conglomerates have stratifications which dip south at similar angles. The stratifications can be traced laterally for 10 m

The remaining 75 percent of the stratified conglomerates have subparallel stratifications which can be traced laterally for 15 m

The conglomerates with low angle stratifications are frequently truncated by conglomerates with subparallel stratification. Pebble zones are common (Fig. 13, p. 32).

Alternations of medium conglomerate and coarse sandstone occur at two horizons in the sequence. The sandstones range from 4 to 50 cm thick and have coarse conglomerate clasts up to 14 cm long within them. The sandstones vary in thickness laterally, some become channellized, others are eroded by the overlying conglomerates. One channellized sandstone 20 cm thick thins eastward and its base becomes gradational in that direction. The sandstones cannot be traced laterally for more than 30 m. Most of the sandstones are graded but in each alternation the bottom sandstones are as coarse as those at the top. Coarse grained subparallel layers which average 3 cm in thickness occur in some sandstones. These layers are persistent laterally, some thicken to form medium conglomerates. The sandstones which appear to be bimodal



Figure 28: Stratifications in a medium conglomerate with scattered pebbles and boulders.

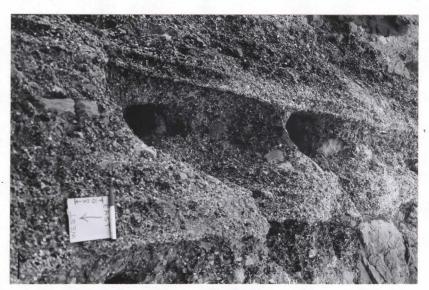


Figure 29: Low angle stratifications in a medium conglomerate with scattered pebbles and boulders.

have angular to well rounded grains.

All the unstratified conglomerates grade into thin layers of fine conglomerate which are less than 10 cm thick.

Composition

The composition of the facies is summarized in Table 10, p. 54 (also Figs. 14-17 Map 2, Appendix 4). The low percentage of matrix suggests the facies is grain supported. The pebbles and boulders are similar to those in the coarse conglomerates.

Roundness and Elongation

The average elongation indexes (the averages of the elongation measurements at each locality) range from 2.0 to 2.6 (Appendix 2, Map 3). They do not show any significant change vertically throughout the section. No significant change occurs vertically within any bed.

The average roundness values range from 0.39 (Bed 62) to 0.50 (Bed 51) (Dobkins and Folk, 1970), but no systematic change occurs up section (Map 3, Appendix 1). The grains vary from well rounded to angular (Fig. 30).

Imbrication

Measurements of the easterly dipping imbrication are summarized in Table 11, p. 56 (also Appendix 5, Figs. 19 and 20, Map 4). The imbricate grains are separated by fine conglomerate, coarse sand and matrix (Fig. 31). Within any one bed, imbrication is laterally variable. The variations are summarized in Table 12, p. 56. Bed 36 has two reversals of imbrication, that is, grains with their long axes dipping westward. The angles of imbrication of 1 and 2 degrees (long axes dipping westward)

Constituent	Range (%) Throughout The Section (Bed And Sequence Nos. In Parenthesis)	Range (%) Within The Sequences (Bed And Sequence Nos. In Parenthesis)		
Sandstone	3(14,1) - 25(26,2)	3(14,1) - 6(7,1)		
		6(49,2) - 25(26,2)		
		6(57,3) - 15(65,3)		
Quartz	2(65,3) - 40(7,1)	17(14,1) - 40(7,1)		
		4(26,2) - 21(23,2)		
		2(65,3) - 8(57,3)		
Micrite	21(7,1) - 42(51,3)	21(7,1) - 28(14,1)		
		26(23,2) - 35(26,2)		
		28(52,3) - 42(51,3)		
Matrix	12(52,3) - 35(14,1)	24(7,1) - 35(14,1)		
		17(26,2) - 31(49,2)		
		12(52,3) - 28(57,3)		

TABLE 10. SUMMARY OF THE COMPOSITION OF THE FACIES⁴ (Medium Conglomerate)

 4 All those grains with apparent a-axes less than 1 mm were classed as matrix.



Figure 30: Range of roundness in a medium conglomerate with scattered pebbles and boulders.



Figure 31: Imbricated medium conglomerate with scattered pebbles and boulders. The figures on the tape decrease eastwards.

(Medium Conglomerate) No. Of Grains Range Of Imbrication In Degrees (Bed And Sequence Nos. In Parenthesis)					
50 and 25	9(38,2) - 20(4,1)				

TABLE 11. SUMMARY OF THE IMBRICATION MEASUREMENTS

are recorded at the same outcrop as an angle of 4 degrees (long axes dipping eastward).

No. Of Grains	Bed No.	Range Of Imbrication In Degrees (Distance Between Localities Shown In Metres)
50 and 25	26	11-14 (1.0 m)
	4	12-20 (0.75 m)
100	58	8-11 (0.75 m)
	38	7-12 (146 m)

TABLE 12. SUMMARY OF WITHIN BED VARIATIONS OF IMBRICATION (Medium Conglomerate)

Fine Conglomerate with Scattered Pebbles and Boulders

General Features

This facies comprises about 5 percent of the sequence (Fig. 32). The types of occurrence of the facies are shown in Table 13, p. 58. The average bed thickness is about 90 cm with some beds ranging from 10 to 160 cm (Map 1). Some of the conglomerates are channellized with channels up to 10 cm deep. This facies does not display any slump

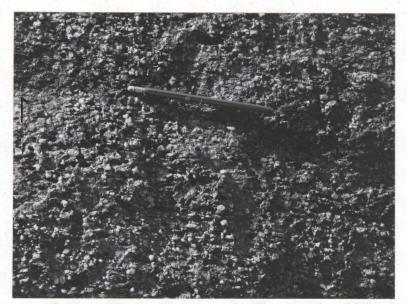


Figure 32: Outcrop of Fine Conglomerate with Scattered Pebbles and Boulders.



Figure 33: Wavy stratifications associated with fine conglomerates and coarse sandstones.

Types Of Occurrences Of The Facies	Whole Bed Stratified	No Structure At All
Individual Beds Of		
Fine Conglomerate		1
Fine Conglomerate		
Grading To Coarse		
Sandstone	3	3
Fine Conglomerate Graded		
From Medium Conglomerate	1	et 1999

TABLE 13. TYPES OF OCCURRENCE OF FACIES (Fine Conglomerate)

or load structures.

The largest clast (140 cm) and the maximum D/10 (76 cm) are in the fining upward sequence (Bed 16, Map 3, Appendix 3). The coarse conglomerate clasts are scattered throughout the bed. There is no relationship between clast size and clast lithology. The lithologies of the coarse conglomerate clasts are similar to those in the other facies.

Internal Structures

Fifty percent of all the conglomerates are stratified. The stratifications are picked out by increases in the percentage of quartz. Single low angle stratifications with an apparent dip of up to 4° east can be traced for 6 m. Subparallel stratification can be traced laterally for 10 m. Wavy stratifications with amplitude up to 10 cm and wave length up to 1 m can be traced over 10 m. On any one stratification neither the amplitude nor the wave length are constant. The stratifications are up to 30 cm thick (Fig. 33). Pebble zones are common.

At two horizons the conglomerates of this facies alternate rapidly with coarse sandstones which are up to 50 cm thick. Some of the sandstones are eroded by the overlying conglomerates. In alternations 46 and 49 the basal sandstones are coarser than the upper sandstones. Coarse grained layers composed of quartz and feldspar (grains with an apparent a-axis 0.2-0.7 cm long) are common in the sandstones. The layers, which are persistent laterally, are not graded (Fig. 34). The thinnest sandstones (< 5 cm) are developed at the top of wavy and subparallel stratifications (Fig. 33). Clasts of fine conglomerate size in the sandstones generally have an orientation subparallel to bedding. The sandstones, which appear bimodal, have well rounded to angular grains.

Composition

The composition of one bed (Bed 18, Map 2) in the first fining upward sequence is summarized in Table 14 (Appendix 4).

TABLE 14. SUMMARY OF THE COMPOSITION OF BED 18⁵

the state of the s	•
Constituent	Percentage
Sandstone	1
Quartz	30
Micrite	21
Matrix	41

⁵All those grains with apparent a-axes less than 1 mm were classed a matrix.

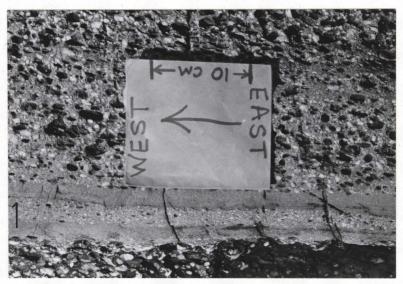


Figure 34: Coarse grained layers in the sandstones associated with the fine conglomerates.

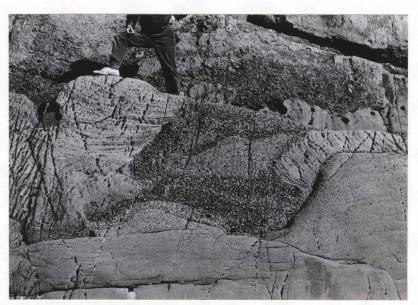


Figure 35: Trough cross-stratification in a fine conglomerate with scattered pebbles and boulders.

Elongation, Imbrication and Roundness

The average elongation index for one bed is 2.4 (Map 3). The elongation ranges from 1.00 to 6.5 (Appendix 2) but no consistent change occurs vertically. Imbrication is not developed but grains subparallel to the stratification are common. The grains are separated by coarse sand and matrix. The grains are well rounded to angular.

Palaeocurrent Measurements

Two measurements of trough cross-stratification (Bed 50, Map 1, Fig. 35) give palaeocurrent directions of 163 and 193 degrees. These directions compare favourably with those presented by Mathey (1970) but differ considerably from the palaeocurrent directions based on grain orientations presented in this study.

Coarse Sandstones

General Features

This facies comprises about 18 percent of the sequence. The types of occurrence of the facies are shown in Table 15. The average bed

TABLE 15. TYPES OF OCCURRENCE OF FACIES

(Coarse Sandstones)	
Coarse Sandstones With Sharp Bases	6
Coarse Sandstones Which Grade Up Rapidly From Conglomerates	10
Coarse Sandstones Which Grade Up Gradually From Conglomerates	7

thickness is about 60 cm, with some beds ranging from 10 to 160 cm. At

the base of those sandstones which grade up rapidly from the conglomerates below, there is a transitional zone, marked by a coarse sandstone with clasts of fine conglomerate size. The transitional zone is 3-4 cm thick. Those sandstones which grade up gradually from the conglomerates have similar transitional zones, up to 25 cm thick. In 50 percent of the sandstones isolated clasts of fine conglomerate size occur up to 10 cm above the transitional zones. The sandstones with sharp bases have channels or scours up to 60 cm deep (Bed 27). In some places, injected dikes of sandstone penetrate upward into the overlying conglomerate for distances up to 130 cm. They occur near coarse conglomerate clasts. The texture of the overlying conglomerate is disrupted near the injection structure (Fig. 36). Pebbles and boulders, up to 36 cm long, occur rarely in this facies. In places boulders are associated with load structures which occur nearby. In other cases, the pebbles and boulders rest in the centre of the sandstones and in the overlying conglomerate there is a complete absence of boulders.

In the first and second fining upward sequences the sandstones are eroded by the overlying conglomerates. The sandstones are difficult to trace because of faulting. In the third sequence the sandstones are traced without difficulty.

Internal Structures

Twenty-five percent of the sandstones display dish structure (Fig. 37). The structure occurs in zones up to 45 cm thick, some of which are diffuse but others are sharp and can be traced for 5-6 m laterally. The dishes are up to 10 cm wide and 1 cm deep. The form of the dish structure appeared to remain relatively constant when the zones



Figure 36: Disrupted conglomerate texture near a sandstone injection structure.



Figure 37: Dish structure in coarse sandstone.

were traced laterally and vertically. The structure occurs below subparallel laminations. Sixty percent of the sandstones have subparallel lamination (lamination < 2 mm thick) in zones up to 45 cm thick which can be traced laterally for up to 15 m. Generally the zones extend to the top of the bed but in some cases the subparallel lamination disappears near the top.

Fifty percent of the sandstones exhibit low angle lamination. The laminations, which are less than 5 mm thick, have an average apparent dip of 8 degrees west and single sets (generally less than 10 cm thick) can be traced laterally for 6 m. The tops of the laminations are always truncated, either by the overlying conglomerate or by subparallel lamination. In some beds the low angle lamination overlies the subparallel lamination.

Twenty-five percent of the sandstones display trough cross-bedding, but only one bed gives a reliable palaeocurrent direction. This indicates a current flowing from the northwest (240°). The cross-bedding, which occurs in single sets up to 1 m wide and 30 cm thick, has an average apparent dip of 20 degrees, and the tops of the sets are generally truncated. Within some sets the apparent angle of dip decreases upwards until it becomes subparallel. The trough cross-bedding is frequently coarser grained than the material below it.

None of the sandstones display convolute lamination, but zones of coarse grained lamination (laminations > 2 mm but less than 1 cm thick) occur in 90 percent of the beds. The laminations, which do not appear to be graded, are composed of quartz and feldspars (grains with an apparent a-axis 0.1 cm-0.5 cm). The laminations have sharp bases and tops. The

zones, which are often associated with amalgamated beds, are up to 20 cm thick and can be traced laterally for 5 m. In the second fining upward sequence two sandstone beds amalgamate eastwards (Fig. 38). Fine conglomerate grains are scattered above the junction of the two beds.

Composition

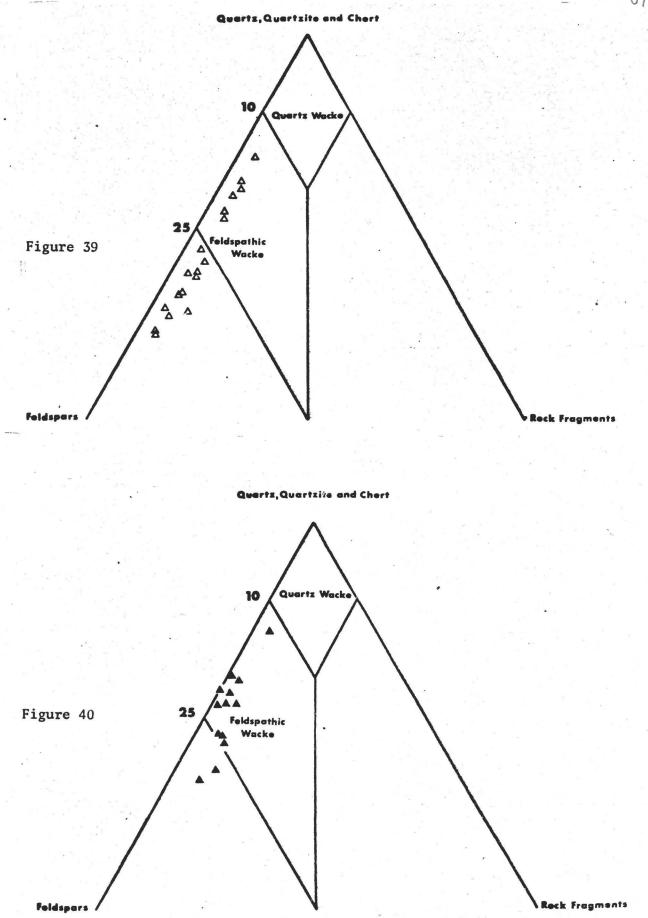
The composition of the sandstones does not vary a great deal (Fig. 39, Appendix 4). The compositions presented here compare well with those presented by Mathey (Fig. 40) (Mathey, 1970). The most common constituents are quartz and feldspar. The feldspars include orthoclase, perthite, microcline and plagioclase. Extremely small percentages of biotite, muscovite, sedimentary rock fragments, zircon and apatite are present. The grains range from well rounded to angular. The matrix is chloritic and very fine quartz and the cements are silica and calcite.



Figure 38: Truncated fine conglomerate which fills a scour in a coarse sandstone.

Figure 39: Composition of the Coarse Sandstones. (Classification after Williams et al., 1954.).

Figure 40: Composition of the Coarse Sandstones. (From Mathey, 1970. Classification after Williams et al., 1954.).



Orientations

It was necessary to compare any palaeocurrent data with that which had been previously collected in the area. The restriction of the measurement of grains greater than 3 cm long was not applied to the orientation study because on the bases of the beds studied most of the grains (over 80 percent) were longer than 3 cm and on the tops most were shorter than 3 cm

The grand vector mean calculated from 6897 vectors was 236 degrees (standard deviation 15 degrees) (Appendix 6). Sixty-nine vector means were calculated and they varied from 220 degrees (Bed 23) to 256 degrees (Bed 29). Fourty-one of the vector means were measured on the tops of beds, twenty were measured on bases and eight were measured at different horizons within certain beds (Map 4). The grain orientation information is summarized in Table 16, p. 69. The four vector means which differ from the grand vector mean by more than one standard deviation are shown on Table 17, p. 70.

No appreciable change in the vector means appears to occur vertically throughout the section or laterally within a single bed. The variation of the vector means in each sequence is summarized in Table 18, p. 70. Lateral differences in orientation (within a single bed) appear to be random. They are summarized in Table 19, p. 71. Mathey (1970) observed nine palaeocurrent indicators in Horizon 3, which are shown in Table 20, p. 72 (see also Table 2). These palaeocurrent directions differ significantly from the preferred grain orientations presented here. The strike of the rocks in Horizon 3 is 233 degrees and this differs by only 3 degrees from the grand vector mean of 236 degrees. The palaeocurrent

TABLE 16. COMPARISON OF VECTOR MEANS WITH GRAND VECTOR MEAN

Facies	Number Of Vector Means Which Differ From The Grand Vector Mean < 1 Standard Deviation	Number Of Vector Means Which Differ From The Grand Vector Mean > 1 Standard Deviation			
Coarse Conglomerate					
Poorly Sorted	2	2			
Coarse Conglomerate Well Sorted	10	1			
Medium Conglomerate With Scattered					
Pebbles And Boulders	48	. 1			
Fine Conglomerate With Scattered					
Pebbles And Boulders	5				

TABLE 17. VECTOR MEANS WHICH DIFFER FROM GRAND VECTOR

Facies	Position In Bed	Vector Mean	
Coarse Conglomerate, Well Sorted	Bottom	220	
Coarse Conglomerate, Poorly Sorted	Within Bed	253	
Coarse Conglomerate, Poorly Sorted	Within Bed	256	
Medium Conglomerate With Scattered Pebbles And Boulders	Тор	252	
	Coarse Conglomerate, Well Sorted Coarse Conglomerate, Poorly Sorted Coarse Conglomerate, Poorly Sorted Medium Conglomerate With	In Bed Coarse Conglomerate, Well Bottom Sorted Bottom Coarse Conglomerate, Poorly Within Sorted Bed Coarse Conglomerate, Poorly Within Sorted Bed Medium Conglomerate With Top	

MEAN BY MORE THAN ONE STANDARD DEVIATION

directions given by Mathey (1970) are approximately transverse to the strike. The imbrication studies show a consistent easterly inclination

TABLE 18. VARIATION OF VECTOR

MEANS IN EACH SEQUENCE

Sequence Number	Range (Degrees) Of Vector Means (Bed Number In Parenthesis)
1	230(21) - 244(16)
2	220(23) - 256(29)
3	221(61) - 247(72)

of elongate grains. This inclination and the grand vector mean of 236 degrees are interpreted as indicating currents flowing toward the

Bed Number (Sequence In Parenthesis)	Range (Degrees) Of Vector Mean	Position In Bed	Distance In Metres Between Vector Means
21(2)	230 - 235	Тор	2.0
49(2)	233 - 252	Тор	3.5
23(2)	220 - 223	Bottom	1.0
42(2)	223 - 238	Bottom	18.0

TABLE	19.	VARIATION	OF	VECTOR	MEANS	WITHIN	BEDS

	Current Direction And Sense Known			Current Direction Known		
Cross-Stratification Channels			Chan	nels		
6	200		160	125		
	185	180	150	020		
			150	015		

TABLE 20. PALAEOCURRENT MEASUREMENTS FOR HORIZON 3

southwest parallel with the present day tectonic axes.

CHAPTER 4

INTERPRETATION

The interpretation of the conglomerate sandstone sequence at St. Simon can be approached by accounting for the most prominent and problematic features of the deposit. The most important features are; the facies distribution, the normal and inverse grading, the imbrication and orientation, and the stratification. Other features include, the range of roundness and elongation, the range of the D/10 values and the occurrence of the thick massive sandstones.

Facies Distribution

Figures 4-7 show that there are few gradational contacts between the facies. Coarse conglomerates generally occur as beds of coarse conglomerate, and do not pass up into thick developments of medium and fine conglomerates. This feature may be related 1, to the grain size distribution of the sediment which was available for deposition; 2, to the mechanism of deposition; or 3, to erosion which occurred after deposition. The occurrence of channels indicates that some erosion occurred after the deposition of the conglomerates. However, it seems unlikely that erosion would result in the consistent development of only 10-15 cm of fine conglomerate at the top of most coarse conglomerate beds. Also, if a thin development of fine conglomerate were formed at the top of a bed, it is difficult to envisage the depositing flow not forming thicker layers of

fine conglomerate if fine conglomerate were available. Thick gradational sequences between the facies are absent because there was a lack of certain grain sizes in the original grain size distribution available for deposition.

Normal and Inverse Grading

The types of grading have been outlined in the facies description. Some of the normal grading in the conglomerates is shown up by a decrease of the maximum grain size upwards through the bed. This type of grading has been called "coarse-tail grading" by Middleton (1967) and he described it (Middleton, 1962) in turbidites of the Ordovician Normanskill and Charny Formations. Middleton (1967) produced the coarse-tail grading in experimental studies on high concentration turbidity currents (approximately 40 percent by volume). Other features of the conglomerates also suggest that the depositing flows were of high concentration. Inverse grading is common in the conglomerate facies. The inverse grading is confined to the bases of the beds. Above the zone of inverse grading, normal grading occurs. Bagnold (1954) suggests that the transfer of momentum, from grain to grain within a high concentration flow, supports the grains throughout the flow. Bagnold suggests that the dispersive force acting upon a particle is proportional to the rate of shear acting on the flow. For the flows which deposited the conglomerates at St. Simon, shear could have been the result of downslope movement of material under the influence of gravity. At the top of a flow, interaction with the overlying fluid medium and an increase in the dispersive pressure might have initiated turbulence. Continued downslope flow could have

resulted in the turbulence spreading down through the flow. Normal grading would result from the settling out of the grains from the turbulent flow. Consequently, those beds which do not display any grading are probably the result of those flows which did not become turbulent.

If the turbulence did not extend throughout the flow, the basal section of the flow would be influenced by the dispersive pressure existing there. If the dispersive pressure were sufficiently high, grains would have moved with some degree of independence. Middleton (1970) has suggested that inverse grading develops because "the smaller particles tend to fall into the spaces between the larger particles and thus displace the larger particles toward the surface". Such a mechanism would operate in a flow where the dispersive pressure allowed some independent movement of the grains, producing inverse grading. The imbrication and orientation also suggest that dispersive pressures operated when the conglomerates were formed.

Those beds which display inverse and normal grading and those beds with inversely graded areas containing boulders, suggest that turbulent conditions existed at some stage in the flow. A change in the flow, from turbulent conditions to conditions in which dispersive pressures were important would account for the formation of inverse grading (by using Middleton's idea that the smaller particles displace the larger particles towards the surface). The time period over which the dispersive pressure operated must have been a short one because the inverse grading does not extend throughout the bed. Middleton (1967) suggests that a decrease in the shear stress in a flow will result in an increase in the viscosity. In the flows which deposited the conglomerates at

St. Simon a decrease in the shear stress would have been the result of the end of downslope movement of the material under the influence of gravity. A rapidly progressive increase in the viscosity upwards from the base would have resulted in the preservation of the grading within the bed. Rapid conditions of deposition and consolidation must be invoked to preserve the grading displayed in the conglomerates.

Imbrication

Some clues for the interpretation of these deposits are also offered by the imbrication. All the facies show imbrication, some to a greater extent than others. Imbrication can be the result of several processes. It has been observed in streams and flumes (Fahnestock and Haushild, 1962) that deposition of a pebble on a sand bed will result in erosion upstream of the pebble, so that the pebble will fall into the scour pocket created. However, this process will not account for the imbrication observed in the conglomerates because not all the conglomerates lie upon sand beds.

Pebbles moved as bed load material in flows become imbricated when they come to rest behind a stationary pebble (Johansson, 1963). The preferred orientation of the long axes is statistically perpendicular to the direction of the flow (Rust, 1972). In the present study the long axes were found to be parallel to, and not transverse to the flow direction suggested by the imbrication. A unimodal distribution of long axes occurs and hence it is unlikely that the pebbles rolled as bed load material.

It has been suggested that dispersive pressures operated during

the formation of the conglomerates. Under these conditions most of the grains will not be in contact with the base of the bed, that is, they will not move as bed load material. Experimental studies on the movement of coarse material in suspension are hindered by the scale of the experiment required. However, some theoretical considerations of the movement of ellipsoidal particles in suspension have been made. Jeffery (1922) considered the movement of ellipsoidal particles in a viscous fluid. He showed that, provided the velocity of the flow were small, the ellipsoids were carried forward rotating in orbits about their centres (Fig. 41).

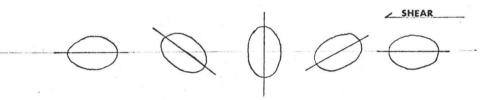


Figure 41. The movement and rotation of an ellipsoidal particle in a shearing medium. (Modified from Glen, Donner and West, 1953).

The rate of rotation varied with the position of the long axes of the pebbles, being at a minimum when the long axes of the pebbles were nearly parallel to the direction of flow. By adopting Jeffery's theory for the rotation of ellipsoids in a plane parallel to the flow, we can account for the imbrication displayed by the conglomerates. Jeffery's theoretical considerations have been supplemented by Taylor (1923). Taylor found that ellipsoids in a viscous fluid under shear tended to assume a transverse position in the flow only if the particles had been in the flow for a long time. Few transverse orientations are recorded in the conglomerate deposits. This might indicate that the time period during which the imbrication and orientation developed was a short one.

Glen, Donner and West (1953) use Jeffery's theory and consider the effect of collision between the particles in the flow. They suggest that, "Collisions may aid the stones to move from orbit to orbit towards the orbit of minimum energy, and in this way could favour the transverse orientation". Manley, Arlov and Mason (1955) also suggest that a high concentration of particles in the flow will lead to the development of a transverse orientation. This transverse orientation would not be compatible with the parallel orientation associated with the imbrication in this study.

Bagnold's theory for the dispersive pressures operating in high concentration flows has been considered for the formation of the inverse grading. Rees (1968) utilizes Bagnold's theory and considers the movement of ellipsoids being sheared in a fluid. He considers the movement of particles in layers, with collisions occurring between particles in adjacent layers and suggests (Rees, 1968) that, "If the collisions are elastic the force in each will be perpendicular to the tangent plane common to the two colliding particles." (Fig. 42).



Figure 42. Collisions between particles in layers with fractional overlap. (Modified from Rees, 1968).

In this way a tilt of the long axes of the particles occurs. This tilt can be interpreted as imbrication. Rees favours collisions which result in the long axes of the particles being parallel to the flow, as in the present study. Thus the imbrication seen in the conglomerates could be produced at a late stage in deposition when dispersive pressures are more important than turbulent conditions of flow.

If dispersive pressures were sufficiently high in a flow, it is not unlikely that the grains would be free to rotate during transportation. It is possible that the movement of the grains would be similar to the movement outlined by Jeffery (1923). Lindsay (1968), utilizing Jeffery's theory, demonstrated by computer simulation that strong long axis fabrics may develop in high concentration flows in a short period of time. In any one situation, the rate at which the fabric develops depends upon the viscosity. The development of the fabric is cyclical. Initially the clasts are in a random orientation. The clasts become oriented with their long axes parallel to the direction of flow and dipping steeply upstream. The angle of plunge of the long axes decreases as the fabric develops. Maximum development of the fabric occurs when the long axes are horizontal and parallel to the direction of the flow. The fabric begins to degenerate when the long axes dip downstream. The strength of a fabric within a flow is related to the time at which the cycle of development is stopped. The cyclical development of the fabric can account for the orientation, imbrication and the reversals of imbrication recorded in the facies.

It was noted that the imbrication died out at the top of many beds. According to Lindsay (1968) this might represent the maximum of

the cycle of development of the fabric. However, if the turbulent conditions which resulted in the formation of grading ceased rapidly and the shearing which produced the imbrication and orientation decreased in intensity upwards through the bed, little imbrication would form at the top of the bed. To preserve the imbrication rapid consolidation of the bed is necessary as has been suggested for the grading.

Channels and Stratification

Most beds have plane to undulating bases but a few are channellized. The plane and undulating bases are probably a function of the fact that the direction of strike and the presumed palaeocurrent direction are the same, hence cross sections of channels would tend not to be exposed. Those beds which exhibit channelling were deposited by currents which did not flow parallel to the present tectonic axes. Some of the medium conglomerates exhibit arcuate stratifications which dip east, west and south. These stratifications are the result of the infilling of scours excavated by currents which flowed in different directions.

Subparallel and low angle stratifications with apparent dip up to 14 degrees occur in some conglomerates. The low angle stratification (p. 49) occurs either as a single stratum or in sets. The stratifications are graded, the finest material occurring at the top. The flows which deposited the stratifications must have been turbulent because many stratifications are normally graded.

Range of Roundness and Elongation

The roundness values range from .38 to .50. The value of .38

corresponds to the mean value of roundness of pebbles collected from streams on Tahiti-Nui, the value .50 is close to the mean value of roundness of those pebbles collected upon beaches on Tahiti-Nui (.55) (Dobkins and Folk, 1970). However, the limitations of using this method for consolidated samples may introduce a significant error in the roundness value of a pebble. This method is of more use on outcrops where the pebbles can be removed intact from the rock.

No consistent vertical change of elongation was noted within the beds studied. This suggests that the flows were not capable of effecting any sorting of the grains according to their elongation. Middleton (1962) also noted that currents which deposited Normanskill and Charny sandstones did not effect a strong shape-sorting. The conglomerates are poorly sorted and this suggests that deposition may have been rapid.

Range of the D/10

No appreciable change of the D/10 was noted throughout the sequence. Because of the outcrop length, it is difficult to make any comment as to whether the competency of the flows decreased in the presumed downcurrent direction. Perhaps later studies carried out over a wider area will show a systematic change of the D/10.

One large isolated boulder in Bed 16B (Map 1) gives some idea of the mechanism of transportation. This boulder has a crack down to its centre which is filled with coarse sand, identical to that in the bed above. The sides of the crack are lined with a laminated layer of calcite about one inch thick. It is most probable that this laminated calcite existed in the boulder before it was transported by the current.

During transportation and deposition intergranular collisions and collisions with the substratum were sufficient to crack the boulder along the calcite layer but not to break the boulder. The boulder came to rest with the crack opening upward, and the crack was filled with sand during deposition of the next bed.

Slumping

There is no sign of slumping within the entire section, and transportation of the beds by "slumping" is unlikely. There is evidence, indicated by the occurrence of load structures, that the sediments were capable of moving under the influence of loading. The load and associated injection structures are not widespread, but are restricted to the top part of each fining upward sequence and always associated with coarse conglomerate.

Composition of the Conglomerates

The increase in the percentage of quartz upwards through the sequence has been noted by Mathey (1970) at St. Simon and by Hubert et al. (1970) at St. Fabien. Hubert et al. (1970) suggest, "In order to provide an increase in quartz, new sources had to become available".

Coarse Sandstones

The coarse sandstones are associated with the conglomerate in three ways. They either grade gradually or rapidly from the conglomerates or they fill channels cut into the conglomerates. There are some sandstones which in parts are channellized but when they are traced laterally they are seen to grade up from conglomerates. This suggests that the depositing flow was behaving differently at various points, depositing in some areas but eroding in others.

Stauffer (1967) describes sandstones which he believes were formed by flows in which grain interactions were controlling the flow. The sandstones described are similar in many respects to those studied herein. Both types contain outsized clasts, and are poorly graded with relatively sharp bases. They are massive but do exhibit subparallel lamination and dish structure. The sandstones lack such features as flutes, grooves and ripple-drift cross-lamination. Bagnold's theory (1954) of dispersive pressures can be utilized to explain the features of the sandstones. Under the influence of dispersive pressure features such as ripple-drift cross-lamination and parallel lamination would not form. However, if the flow had previously been in a turbulent condition, as is suggested by the grading and imbrication of the conglomerates, some parallel lamination could have formed when the flow was decelerating.

Large scale cross-stratification occurs at the tops of some sandstones. Middleton (1969) has suggested that the cross-stratification may be the result of deposition in scours formed by the entrained fluid along the top of the flow. The subparallel lamination seen near the top of some coarse sandstones may also be formed by this mechanism. However, some sandstones are obviously amalgamated (Fig. 38) and the coarse grained laminations seen in some sandstones may be the result of several amalgamations.

Evolution of the Flow

To understand the evolution of the flows which produced the

conglomerate and sandstone sequence at St. Simon, we must account for the mechanisms which have been invoked to explain the sedimentary structures within the deposits. We must account for the turbulence and dispersive pressures, but this problem is difficult to resolve because experimental work on coarse grain deposits and observations of natural processes involving coarse grained deposits are few. Morgenstern (1967) has given some theoretical consideration to the problem of the development of turbulence within flows. He suggests that excess pore pressures must exist in a cohesionless mass moving downslope before turbulence will be initiated in the mass. The method by which excess pore pressures build up in coarse grain deposits is not clearly understood. Morgenstern (1967) suggests that dissipation of excess pore pressure and a rapid decrease in the slope over which the mass flows would prevent the onset of turbulence. Mixing of the sediment on the surface of the flow with the overlying fluid medium is also important because it relates to the density of the flow. The mixing process must be clarified before the mechanics of high density turbulent flows can be explained.

In the light of our present knowledge, the evolution of the flows which deposited the conglomerate sandstone sequence at St. Simon is envisaged in four stages.

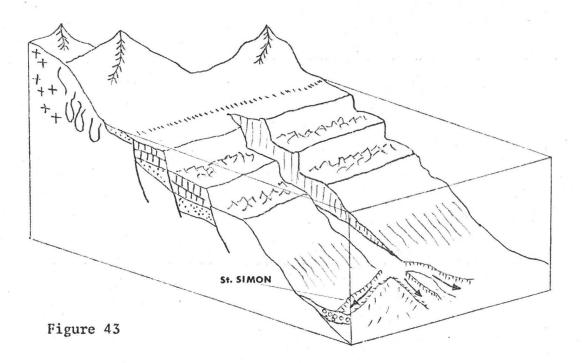
- Initial downslope movement of a sediment under conditions of excess pore pressures.
 - (2) Continued downslope movement under conditions of excess pore pressure initiates turbulence at the top of the flow. Continued movement under turbulent conditions produces normal grading vertically and laterally within the flow.

- (3) As material settles to the base of the flow, dispersive pressures are established and inverse grading, imbrication and orientation are formed.
- (4) Rapid freezing results in the preservation of the internal structures.

Depositional Environment and Palaeogeography

At the present time deep water conglomerate deposits are restricted to submarine fans and submarine channels. It is likely that these environments existed in the past and were major areas of accumulation of conglomerate deposits. For the Québec Group deposition of conglomerates on submarine fans has been suggested by Mathey (1970) at St. Simon and Hubert et al. (1970) St. Fabien and L'Islet. For the St. Simon area, Mathey (1970) envisages a submarine fan complex, built of sandy and conglomeratic material by turbidity currents and slumps (glissements) which flowed transverse to the present tectonic axes. The currents and slumps reached the sea floor by way of a submarine canyon which extended into a shelf sea area.

For the St. Fabien and L'Islet areas, Hubert et al. (1970) propose a similar model to the one suggested by Mathey (1970). "An unstable source area composed of shelf type limestones and sandstones and the Precambrian Shield area provided coarse debris which were channellized and deposited on submarine fans by turbidity currents and other mechanisms.". Hubert et al. (1970) indicate that currents which flowed parallel to and transverse to the present tectonic axes were operative during the formation of the submarine fan complex. Figure 43: Reconstruction of the Cambro-Ordovician palaeogeography in the St. Simon area. The sediments are derived from Palaeozoic limestones (blocks) and sandstones (stippling), and Precambrian metamorphic (folded) and igneous (cross) terranes. (Modified from Hubert et al., 1970.). For the outcrop at St. Simon it is suggested that the conglomerate and sandstone sequence studied exists on that part of the submarine fan complex which was built by currents which flowed mainly parallel to the present tectonic axes (Fig. 43). However the problem to be explained



is why the sequence was deposited parallel to the present tectonic axes. The currents might have flowed along a meandering channel which is parallel to the present tectonic axes or they might have flowed in a distributary channel which is parallel to the present tectonic axes. Both types of channels have been reported on the La Jolla Fan Complex off the coast of California (Normark, 1970). However, the length of the outcrop studied does not enable the distinction between the two channel types to be made.

An alternative theory is that the conglomerates occupy a submarine channel beyond the margins of the fan. A present-day example of such an environment is the Cascadia Channel. The channel runs for about 200 km

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parallel to the present tectonic axes along the coast of Oregon (Griggs and Kulm, 1970). However, an environment similar to the Cascadia Channel would not fit the stratigraphical evidence which has been recorded in the St. Fabien and St. Simon areas. It would not account for the palaeocurrent data recorded by Mathey (1970) in Horizon 3 (Table 20), which indicate that some currents flowed from the north and northwest. Also, it would not account for the lenticular shape of Horizon 3 nor for the thinning of Horizon 3 towards the southeast (Mathey, 1970), that is, in the direction transverse to the present tectonic axes.

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ROUNDNESS MEASUREMENTS

		(1)					(2)					(3)					(7)					(8)					(9)		•
.58	. 25	.22	. 33	. 29	. 38	. 50	.40	. 50	.75	. 33	. 50	. 38	.25	. 50	.57	.19	.43	.70	.57	.35	.67	. 50	. 38	.53	.74	. 40	.67	. 38	.22
. 33	. 19	.21	.25	.33	. 36	. 33	.67	.67	.50	.21	.17	.50	. 50	.33	.57	.43	.27	. 50	. 33	.17	. 20	.75	. 30	.50	.67	. 38	.57'	. 50	.50
.67	. 33	. 38	. 20	.50	.43	.67	.60	.50	.50	.25	.50	. 40	. 50	. 50	.53	. 50	.40	. 33	.75	. 33	.75	.67	. 50	. 33	.44	.29	.67	.\$7	. 80
.56	.40	.56	.57	.57	.43	. 33	. 50	.25	.50	.40	.40	.17	.50	.50	.56	. 58	.40	.29	.60	.60	.13	. 38	. 50	.5.)	.60	. 50	.67	.40	. 33
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.25	.63	.25	. 33	. 38	. 38	.50	. 50	.25	.67	.25	.43	. 38	. 40	. 50	. 50	. 50	.50	.33	.33	.13	.63	. 50	.33	.50	.44	. 38	. 33	.50	.50
. 25	. 33	.17	.83	.40	. 20	. 30	. 25	. 38	.75	.25	.67	. 33	. 33	.43	.43	.43	.43	.50	. 30	.25	.19	. 80	. 53	.5.9	.50	. 33	.67	.50	.50
.23	.67	.25	. 50	.38	.40	. 38	.50	. 25	.75	.25	.17	.14	.67	.75	.76	.63	. 33	. 33	. 29	. 50	. 33	. 33	.75	.50	.33	.60	. 33	. 50	.40
.43	. 29	. 50	. 38	.23	.15	.25	. 50	. 50	. 33	. 33	.14	.14	.67	.60	.42	. 33	.50	.50	.17	.17	.20	.50	. 20	.53	.67	.67	. 33	.67	. 50
.40	.59	.40	. 33	. 38	.25	.43	.67	. 50	.50	. 33	. 33	.17	.25	. 40	. 50	. 29	.80	.60	.10	.21	. 50	. 33	. 33	. 5.7	.40	. 50	. 29	. 50	.67
.10	. 30	. 38	.25	.50	.75	.17	. 38	.25	.67	.48	. 20	. 38	. 33	. 50	. 30	. 19	.43	.17	.21	.23	.25	.25	. 50	.50	.23	. 30	.67	.67	.25
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.21	.50	.44	.67	. 33	. 30	. 38	.75	. 50	.50	.25	. 25	. 50	. 33	.60	.33	.44	.17	.50	.33	.25	.17	.50	.60	.50	.50	.50	.50	.57	.50
. 38	.40	. 30	. 60	.67	. 38	.53	.50	.67	1.00	.50	.17	. 33	.67	.25	. 38	. 38	.75	.33	.25	.22	. 25	. 33	. 50	.50	.43	.60	. 50	.50	.67
.43	.\$7	.40	.17	.42	. 33	.43	. 38	.40	.75	.40	. 50	.75	.50	.25	.33	.25	.50	.21	.55	. 30	.22	.50	.25	.73	.40	.50	.75	.50	.50
. 29	.33	.25	.13	.70	.40	.50	.67	. 50	.40	.75	.25	. 50	.25	.25	.42	.40	. 38	. 50	.40	.25	. 38	. 33	.75	.75	.33	.57	.67	.80	.17
.38	. 25	.50	.50	.67	.75	.50	. 50	. 50	. 38	.20	. 30	.50	.50	.50	.80	.13	.40	.71	.33	.21	.38	.17	.67	. 20	.38	. 38	. 33	. 33	. 50
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.38	.40	. 50	. 30	.50	.60	. 38	.67	. 50	.33	.17	. 33	.17	. 33	.20	. 26	.42	. 38	.40	.14	.60	.25	.33	.75	.50	.40	. 50	.50	.67	.60
.67	. 33	. 50	. 33	.33	.67	.71	. 40	. 50	.50	. 25	.75	.67	.25	. 38	. 38	.17	. 29	.50	.67	.33	. 38	. 29	.50	.67	.43	.50	.67	.75	.67
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. 38	.50	. 50	:67	.33	. 33	.17	.43	.50	. 10	.40	. 33	. 38	.50	. 33	.17	.33	.67	.67	.33	.21	. 30	. 38	. 30	.50					2
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	50 .	. 50	. 50	.75	.33	.60	.25	.43	.50	. 30	.17	.75	. 20	. 33	.17	. 30	.60	. 30	.75	.31	.80	.14	.50	. 33	.25	
.1	19	. 25	.60	.50	. 50	. 29	.67	. 50	.67	.25	.14	. 29	.60	.50	.50	.50	.25	.25	.20	.33	.50	.50	. 33	.60	. 33	
.1	75 .	.25	.57	.57	.75	.14	.17	. 38	.50	.25	.83	. 60	. 50	. 33	.20	.57	. 30	. 75	.60	.63	. 20	.40	. 33	.50	.61	
.1	75	.67	.67	.75	.67	.43	.40	. 53	.33	.50	.62	. 38	. 38	.50	.40	.50	.50	.55	.33	.50	. 38	. 38	. 29	.40	.43	
.5	50 .	. 25	. 33	.75	.67	. 20	.33	. 50	.50	.67	.50	.57	.40	.40	. 38	.21	.33	.25	.50	.50	.20	. 75	.44	.50	. 5-7	
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.:	33	.33	. 50	.57	.40	.14	.40	.40	.67	.25	.75	.75	.57	.43	.40	. 29	.67	.25	.50	.25	.50	.24	.60	. 30	.25	
	57	. 29	.40	.75	. 38	.25	.11	.75	.50	.25	.71	.17	.25	.40	.50	. 33	. 38	.42	.25	. 38	.67	.25	. 29	.60	.29	
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	50	.40	.50	.67	.67	. 50	.33	.50	. 38	. 33	.25	.25	.50	. 50	.25	.60	.38	.43	.63	.17	.67	.50	. 38	.40	.25	
	75	.50	.75	. 38	.50	. 33	.21	.50	.50	. 33	.43	. 30	.50	.60	.50	.50		.47	.33	.67	.50	.25	. 29	.21	.25	
	30	. 33	.50	.13	.50	.25	.57	.67	.25	.75	.33	.40	.25	. 30	.33	.14	. 38	.38	. 50	. 33	.67	.75	. 38	.27	. 50	
	17	.43	.50	. 80	. 38	. 25	.67	.50	. 29	. 33	.50	. 33	.40	.44	.25	. 80	. 33	.71	.21	.40	.27	.25	. 22	. 40	.67	
	17	. 22	. 50	.50	.50	.25	.38	.50	. 83	.40	.75	. 50	.40	.75	. 50	. 33	.60	. 30	.50	.25	.33	. 38	.40	.50	.71	
	38	. 19	. 50	.75	.67	.17	. 33	.25	.13	.40	.22	.50	. 50	. 50	.50	.25	.44	.42	.60	.28	.38	.20	. 25	. 50	.50	
	67	.40	. 33	.33	.50	.17	.33	.33	.33	.50	.67	.33	. 80	.57	.44	.83	.33	.17	.33	. 38	.25	.75	.25	.46	.43	

Study number in parenthesis.

ELONGATION MEASUREMENTS

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1.00	1.00 1.83 1.67 1.86 3.75 1.20 1.50 1.40 2.27 2.25 1.40 1.11	5.86 2.36 4.75 2.20 1.50 2.00 2.00 2.00 2.01 2.21 1.40 4.57 6.67 2.39 7.00 7.00
1.58	2.00 1.75 1.91 3.00 2.20 1.83	2.40 2.80 2.11 1.94 3.40 2.65 2.87 2.13 3.60 2.00 2.00
2.14	3.00 1.30 1.67 1.50 1.43 1.60 1.67 1.20 2.00 2.64 3.00 2.00	1.54 6,00 2.57 1.33 1.33 1.29 1.33 1.29 1.30 2.03 1.33 2.50 1.33 2.50 1.33 2.52 2.62 1.22 1.22 1.85
(5)	2.25 4.50 2.60 1.80 1.56 1.80 2.60 3.00 1.50 2.00	2.43 2.27 2.65 3.23 2.75 1.22 1.53 3.82 2.83 3.1.53 1.23 1.53 2.25 2.25 2.25
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2.06	1.65 1.11 1.37 1.23 2.22 2.91 1.86 2.89 3.38	3.00 4.00 1.33 2.33 2.00 1.67 3.50 4.25 4.00 3.20 2.33 1.43 4.33 4.20 3.25 1.60 1.83
1.67	2.48 1.53 2.67 2.44 2.10 2.07 2.24 4.71 1.64 1.73	2.40 1.67 1.67 4.50 1.50 5.50 6.50 1.20 3.00 2.43 2.175 1.22 2.50 2.43 2.175 1.22 2.52 2.25
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5.00	1.41 1.10 3.00 2.10 2.13 1.32 3.00 1.29 1.00 2.77 1.09 2.80	
(6) 2.40 1.43	2.50 4.20 2.08 2.51 2.26 3.75 2.29	
2.14	1.58 1.53 1.44 3.68 2.36 2.00 3.10 4.29 1.76 1.19 1.87	
2.80	1.02 2.33 3.13 2.43 2.50	
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Study number in parenthesis.

		0	0										· · ·			(10)							
3.00 1.18 1.21 2.43 3.75 1.63 2.43 1.60 3.50 2.67 2.50 2.60 1.74	2.93 1.50 4.77 3.35 1.71 1.33 3.14 2.80 1.69 2.01 3.00 3.86 3.00 1.77	1.72 3.18 1.97 2.50 2.25 3.20 1.58 3.33 1.29 2.33 2.00 3.07 2.23	3.08 3.28 1.26 2.55 3.00 1.23 5.50 1.23 5.50 1.23 2.59 1.20 2.59 1.20 2.50 1.81 2.25 5.64 1.36 2.41 2.80 2.41 2.80 2.43 2.33	3.91 1.46 4.71 1.24 5.50 3.10 1.10 2.00	1.93 1.91 2.44 4.05 2.60 1.50 1.50 1.50 1.50 1.50 1.50 2.00 1.60 2.00 1.60 2.00 1.60 2.03 3.67 2.31	(8)								1.40 2.09 1.52 1.51 2.67 2.52 2.59 1.60	2.17 1.85 3.20 1.78 3.20 1.78 1.58 2.45 1.52 2.45 2.45 2.45 2.45 2.52 1.52 2.11 4.25 5.50 1.55 2.52 1.550 2.52 1.550 2.38 2.52 2.38 2.52 1.550 2.52 1.550 2.53 2.53 2.53 2.53 2.54 1.550 2.550 2.550 1.550 2.550 2.550 1.550 2.550 2.550 1.550 2.5500 2.5500 2.5500 2.5500 2.5500 2.5500 2.5500 2.5500 2.5500 2.5500 2.55000 2.5500 2.55000 2.55000 2.55000 2.550000000000	2.57 1.33 3.00 2.35 1.53 1.15 2.22 2.40 1.67 1.62 1.07 1.62 1.07 1.65 1.40 1.23 1.74 1.51	5.00 1.30 2.43 5.60 1.15 1.31 1.85 2.33 2.45 2.34 2.20 2.34 2.23 1.66	4,00 1.57 2.10 2.58 3.37 2.78 2.38 1.98 3.71	1.94 1.65 3.00 1.54 1.41 1.88 1.29 2.26 1.43 2.88				
4.29 1.20 4.55 2.14 2.20 1.10 1.75 1.89 1.94 3.44	1.40 2.00 3.17 2.00 1.40 1.42 1.63 1.48	1.93 1.57 1.47 2.40	3.80 2.00 1.21 1.20 4.27 1.78 1.50 1.95 1.40 1.56 2.57 2.41	4.00 1.25 2.15 1.15 2.50	1.50 2.69 1.37 2.31 1.35	1.71 1.56 1.86 1.64 1.53 3.02 1.73 1.53 3.02 1.78 1.46 5.29 2.38 2.07	2.89 1.70 1.38 1.15 2.60 4.40 1.47 1.94 1.50 3.04 3.10	1.35 1.01 2.00 1.93 2.67 4.40 2.05	2.67 1.64 1.67 2.22 3.94 2.00 1.81	2.43 5.67 2.23 2.61 1.27 2.72	1.87 1.68 2.89 1.20 1.73 3.86 2.43	1.68 2.24 1.17 2.09 1.03			1.36 1.10 1.03 3.00 5.49 1.25 1.60 4.16 1.41 1.63			(11					
			(9	1										1.52 3.64 2.95 2.85 1.63 1.67 1.68 2.07	2.52 1.90 4.70 3.40 2.80 4.11 2.44 4.50 2.53	2.14 2.00 2.23 1.52 1.14 1.11 3.00 1.61 3.25 2.64	1.38 1.91 1.71 2.38 1.95 2.00 3.51 3.98 1.43 1.11 1.41	1.23 2.29 1.33 1.63 1.02 1.54	1.50 1.96 2.43 1.21 1.67 3.00 6.77 4.07 1.49 2.03 2.00	1.92 1.79 2.80 1.89 1.56 3.04 2.30 1.64 1.31	1.59 1.00 5.00 1.42 1.36 1.83 1.02 1.95 4.20 2.47 3.09 1.95 1.39	1.70 2.50 4.91 2.70 1.33 2.89 2.55 4.20 3.88 1.94 2.83 2.10 1.11 6.00	3.12 2.82 1.00 2.60 3.47 1.51 2.90 2.11
1.29 2.20 2.71 2.00 1.78 1.40 1.11 2.25 3.22 2.00 2.14 8.33 2.54 2.54	1.89 2.44 3.67 1.60 2.08 2.19 2.57 4.00 4.25 1.54 1.18 2.13 2.08 2.17	1.67 2.00 2.07 2.92 3.57 6.00 1.69 2.00 2.80 2.24 2.23 2.33 3.50 2.00	2.57 2.89 2.92 1.92 4.38 3.00 1.50 2.10 1.11 3.11 3.82	2.89 3.44 2.67 2.45 1.69 1.50 2.13 1.57 2.80 1.80 3.79 2.23 2.40 1.29	2.07 1.75 1.80 2.60 2.70 2.22 2.67 2.14 3.20 2.91 2.33 5.45 2.29	2.00 3.00 1.78 2.64 1.27 2.36 2.42 1.50 1.17 4.29 2.60	1.71 1.80 1.75							1.80 2.55 3.65 4.89 2.40 1.40	1.71 1.55 2.52 2.54 1.32 1.43	1.38 1.83 3.50 3.00 2.13 2.13	2.04 3.40 1.29 1.58 2.26 2.06	(12) 2.80 2.00 1.56 1.36 1.63 1.50	2.90 1.15 3.73 1.00 1.29 2.18	1.67 2.20 1.97 1.73 3.50 1.75	3.54 1.64 1.42 1.72 1.94 3.44	1.23 2.30 1.34 2.56 1.57 1.42	
	1.27 3.11 2.57			2.85										4.83 2.06 1.17 1.62	2.33 1.25 1.43 3.19 4.33	2.36 2.74 2.62 1.40 1.40 1.35	1.18 1.60 1.27 2.81 2.00 1.63 2.33 5.04	3.67 1.71 2.90 2.29 1.71 2.24 4.48 2.14	2.73 1.79 3.18 4.83	2.33 2.26 3.53 2.14 3.25 3.43 2.60		1.42 2.21 1.29 2.27 1.72	

Study number in parenthesis.

MEASUREMENTS OF THE TEN LARGEST CLASTS

(1)	(2)	(3)	(4)	(5)	(6)	(7)	(8)	(9)	
							(a)	(b)		
26	41	34	32	30	64	50	97	29	110	
20	28	38	36	40	49	70	52	37	125	
21	39	23	21	26	33	70	36	59	82	
18	67	24	18	20	33	33	40	53	74	
17	59	23	23	31	43	33	29	47	74	
21	19	23	20	20	23	33	34	27	60	
19	27	23	21	20	20	45	42	25	66	
22	38	22	23	18	16	36	30	26	62	
18	14.	21	19	16	21	33	39	38	49	
19	18	21	23	16	25	34	32	45	48	

(10)	(1)	1)	(12)	(1	3)	(14)	(15)	(1	6)	(17)	(1	8)
4. I.	(a)	(b)		(a)	(b)			(a)	(b)		(a)	(b)
100	51	59	174	- 36	121	36	115	53	32	. 27	51	71
80	63	41	64	39	64	. 30	70	26	145	23	47	35
63	23	32	56	53	59	26	48	30	66	19	45	37
64	74	44	110	25	72	27	39	32	107	18	41	43
60	44	39	59	32	32	36	39	40	77	17	73	-59
59	27	34	100	32	. 36	20	36	33	32	20	44	20
56	25	22	88	31	37	24	32	36	19	15	44	18
54	24	27	83	30	43	28	32	82	27	16	40	25
51	25	23	80	29	46	22	30	57	.19	14	32	27
58	31	34	64	30	42	20	29	265	14	17	29	22

(Study Number)

a: Limestone

b: Sandstone

(1	.9)	(20)	(21)	(22)	(2	3)	(24)	(25)	(26)	(27)
(a)	(b)				(a)	(b)				
35	60	128	58	59	20	26	49	51	63	24
32	36	98	40	45	19	29	115	49	60	44
28	46	86	52	165	24	22	34	36	74	33
32	54	69	41	72	19	18	52	38	66	36
44	62	59	46	33	28	16	48	43	58	24
31	42	65	41	33	22	22	56	55	60	20
30	29	41	36	33	26	32	32	39	63	21
15	27	54	36	39	24	38	83	37	45	19
15	25	40	35	40	17	27	29	36	54	17
16	24	42	36	43	22	20	32	40	70	18
(28)	(29)	(30)	· (31)	(32)	(33)	(34)	(35)		
			(a)	(b)					×	
46	32	29	25	26	46	55	69	76		
42	41	32	22		- 59	- 50	43	- 68		
34	39	25	33	16	40	43	49	39		
22	35	26	21	33	63	45	59	33		
25	59	20	19	27	34	45	43	29		
- 22	- 38		25	24	96	115	39	27		
23	31	17	16	22	48	52	43	34		
17	36	15	15	15	36	150	40	23		-
19	36	16	29	16	37	54	51	23		
23	47	19	26	23	28	103	39	33		

(Study Number)

a: Limestone

b: Sandstone

COMPOSITION OF CONGLOMERATES AND COARSE SANDSTONES

1							· ·		
Study Number	1	2	3	4	5	6	7	8	9
Bed Number	2	7	14	18	23	26	33	33	35
Matrix	17.2	24.4	35.2	40.6	30.8	17.5	12.4	12.5	34.5
Grey Micrite	26.1	21.2	28.5	20.7	23.2	28.9	22.9	19.8	25.7
Sandy Limestone	19.8	3.8	11.0	3.4	8.0	15.9	21.4	26.0	13.1
Oolitic and Laminated Limestone		2.7	2.2	2.5	trace			trace	8.5
Biomicrite	6.8	trace	trace		3.0	5.9	6.3	7.2	trace
Sandstone	27.0	6.3	3.3	trace	9.3	25.5	21.4	22.7	9.3
Shale	trace			trace	trace	trace	6.0		trace
Quartz	2.4	40.5	17.3	30.2	21.3	4.4	8.2	2.1	7.7
Others	trace	7.2	trace						
Feldspar	trace								

COMPOSITION OF CONGLOMERATES

			i.						
Study Number	10	11	12	13	14	15	16	17	18
Bed Number	36	49	51	52	55	57	63	65	65
Matrix	15.5	31.5	15.2	12.3	27.6	28.6	22.2	11.6	24.2
Grey Micrite	27.4	33.4	37.8	23.1	34.1	35.3	32.8	20.8	34.3
Sandy Limestone	19.8	13.4	20.4	26.3	28.3	13.4	12.6	26.1	11.5
Oolitic and Laminated Limestone		4.3	trace		6.3	4.2	trace	trace	trace
Biomicrite	3.3	trace	4.0	5.6	5.9	3.0	5.0	7.1	7.8
Sandstone	21.8	6.2	15.4	21.4	trace	6.2	16.3	21.3	14.7
Shale	2.6		2.4	6.3	,	trace	5.0	9.0	4.1
Quartz	9.1	10.1	3.4	3.3	trace	8.3	5.2	trace	2.3
Others	trace	trace	trace	1 .	trace	trace	trace		trace
Feldspar	trace		1	trace					trace

COMPOSITION OF CONGLOMERATES

Bed Number	Study Number	Quartz	Feldspars	Matrix	Cement	Others
19	1	69.3	19.0	0	11.7	0
26	2	74.0	10.3	3.0	12.0	1.7
35	3	65.3	18.4	1.0	14.0	1.3
35	4	56.0	19.0	4.3	18.7	2.0
27	5	68.7	16.4	.3	14.3	.3
50	6	52.3	21.0	7.0	16.0	3.7
50	7	65.3	19.7	2.0	12.3	.7
70	8	63.4	17.0	4.3	13.0	2.3
61	9	59.3	20.0	3.3	14.7	2.7
44	10	70.3	16.7	1.0	11.0	1.0
61	11	60.0	27.0	4.0	6.0	3.0
50	12	60.7	20.9	11.0	4.2-	3.2
18	13	65.4	17.0	4.0	12.3	1.3

COMPOSITION OF SANDSTONES

IMBRICATION MEASUREMENTS

	1 1	opeq		PAVIE									ADEDA	TOP	PAVIES
INBOICA	TTON -	5				VIATION	13 05	CREFS.				INSPICATION .	15		STANDA
288	2 * *	273	201	104	792		203	294	200	272	231	295 295	272	281	2#1
240	101	235	304		142	274	25.8	294	248	203	795	787 798	777	704	
299	75 7	269	202		205		273	247	257	280	292	277 364	200	791	273
287	757	284	754	766	745	742	248	244	253	315	324	295 269	273	202	530
277												272 306			
0117000			1100	ONVIER								SUTCROP 184	OPEDA	TOR	LANIES
246	274	14	747	STAND: 284	727	VIATION	397		Sec.			Inusication =	14		STANDA
273	105	295						233	277	313	794	794 301	272	291	771
200		245	272	296	135	277	274	297	232	SBB	788	700 202 777 702	275	260	
ANTCOND	21		1700	-								277 202 294 290	274	244	777
TUDDICA		9	4103			VIATION						794 293	244	244	1.71
297	797	775	777	3.00	244	277	260					OUTCOCP 104	ADEDS	Ton	-
275	797	260	797	176	122		269	221	291	281	272	INDOICATION .	- 1	11.1.4	STANCA
766		\$12.2		140		117	204	//1	Sai	209	300	201 100	297	757	202
	4.5		AT-2	AVIES								774 794	261	213	774
INDOICU		15				VIATION	6 35	ARFES				200	201	~1-	
777	203	299	204	301	273	234	204	200	290	278	302	CUTCOD 204	ADEDA	TOP	DAVIES
770	272	282	321	294	202	279	204	ZRO	276	279	284	THARICATION -	12	1.0-	STANDA
794	* *	E CAL			1.00	214	5.04	240	215	214	200	217 307	204	214	
		1055		DAVIES								271 209	298	201	304
		15				POTATION	4 05	CPEES				794	2.40	2-1	304
797	ACC	201	794	797	707	785	295	293	777	202	312	CUTCOCD 214	COFOA	TOQ	DAVIES
754	200	789	207	287	285	293	290	296	273	256	301	THEOTCATION =	20		STANDA
299				ц.				£ 7.0	2.7	290	201	272 204	205	786	117
Trano	63	1000	AT-2	DAVIES								287 300	299	200	207
	T ! ~ ! =	14				ATTON.	15 05	COFFE				303 297	280	202	302
777	775	756		766	793	336	203	199	247	298	297	. 793 798	2.82	205	203
201	202	797		700	724	255	327	201	207	257	292	304 292		-	6.07
312								e		2	2.72	CUTCOCP 224	CPERA	TCR	DAVIES
Trong	74	-263		AVIES								I"BRICATION .	6		STANDA
VADICA		13		STANDA	on ney	INTION	4 05	REFES				316 251	765	278	797
79.0	726	297	101	292	779	757	294	263	202	287	.755	281 202	755	777	205
794	777	200	293	277	203	283	292	273	297	271	284	777 769	292	105	203
274	291	202	275	270	794	273	249	287	795	249	299	304 273	286	275	272
700	790	280	2.5	285	121	206	240	272	707	291	750	274 279	278	279	111
777	794					1.14		e . c			12-	273 264	251	240	107
117-000		apep!	17-2	DAVIES								783 766	219	206	779
MODICA		16		STANDA	on ney	VOITATION	10 05	OPFES				227 280	252	760	279
797	212	297	740	252	794	292	277	279	294	295	797	257 709	250	771	
275	326	277	275	277	233	305	318	272	273	287	275	OUTCOOD 274	POFDA	TOD	FRAFT
274	267	292	295	293	2=2	286	292	281	315	313	278	THADTCATION =	7	N 12	STANDA
270	767	260	224	375	799	327	296	297	270	280	30 9	257 305	282	297	292
745	2.7.5			- 6			2.4.5	6	2	240	20-	246 252	314	204	302
												296 294	270	206	278
												304 274	797	275	272
Trong	24	2053	17-2	DAVIES								265 263	272	210	747
****		7				POTTATION	10 05					773 765	251	240	202
793	274	261	- 4 4	750	1 24	772	205	303	274	256	274	745 273	269	774	202
777	767	242		220	776	249	108	295	784	200	741	228 291	253	2= 2	270
250	275	274	7 . 1	292	7=3	229	300	307	272	275	264	268 200	251	271	
266	24 =	20%	792	242	=17	271	252	304	235	236	272				
243	7= 4	4		24-		٤٠٠	2.9-	3	5	2.40	216				
UTCOOD		COFO	AT - 2	DAVIES											
VASICA.	T! "" =	. 12				INCITATION.	11 05	CREFS							
221	776	271	776	274	773	26.3	777	297	275	228	224	0117000 775	POFOA	100	PAVIES
787	777	202	274	200	727	224	297	250	236	327	274	THOOICATICH =	1		STANCA
774	205	213		275	202	7=3	300	290	252	745	724	793 753	275	767	757
287	224	307	2.2	202	275	273	296	272	297	276	288	277 250	279	294	273
757	22											272 203	275	279	250
Tronn	114	-057	1777	NAVIES								275 262	371	767	272
10100	TI =	1 =		STANDS	ישר רד	POTATION	8	ROFFS				750 775	275	283	757
797	777	275	2.7	275	775	302	278	201	302	328	224	747 767	266	776	757
202	200	1.50	30=	200	111	2=3	231	321	794	277	221	312 313	220	744	200
770	72%	774		277	777	275	797	222	774	270	7 4 1	257 252	291	793	271
761	72=	771	2-3	770	77=	275	234	200	222	202	725	784 249	240	505	5 300
777	220											OUTCOOD 234	POLOT	TOP	WOLFF
Trono		1050	11-2	DAVIES								INBDICATION =	1	270	STANDA 257
Haster.		12		CTAND: 275	30 .EA	TATICAL		CDEEC				293 250	274		
297	794	207	754	275	221	245	224	249	270	245	269	272 234	279	279	273
250	796	267	777	300	2 - 5	775	291	296	208	300	101	275 269	177	769	272
774	1.1													794	252
כיבידיי		-0523		-AVIES									275	225	257
POTCA	375	14	1.0			1.11.11	17 05					243 267	266	273	289
705		787	3	769	722	769	203	297	205	274	1.1	271 283	291	253	257
272	200	291	722	200	200	205	244	294	202	207	224	244 249	238	203	671
200	272	254	500	762	242	275	301	255	172	243	232	CUTCOCP 74A	ODEDA		DAVIES
279	205	274	200	501	221	503	2==	320	254	277		THROTCATION -	10		STANCA
744		1052										747 777	273	110	273
	74.5			TAVIES								272 271	273	250	
1001CA		9 797	727	CT AN INA	20 759	141174						302 200	289	104	271
775	722			200	7=4	745	742		75 3	255	727	297 235	257	275	
250	202	201	273	4/12	222	171		244	101	309	24?	777 777	200	777	
Trong												104 200	373	257	
Trong				DAVIES								774 767	284	784	
	793	11		STAN, DA		141100			100			701 701	270	794	
unotral	202	275	744	227	747	720	376	272	771	277 233	274	776 763	290	299	
37=		315		214	/	/	5.0	2:1	1.1	213	. 5 4 1	TITCOOP 744	PPEDA		
240												VADICATION -	11	11.00	STANOS
775 260 714				CAVIES								102 200	271	200	
240 240 214		0	200			TATION						271 271	100	204	
775 740 714 714				202	273	300	223	230	245	200	237	100 200	297	207	
249 249 314 17000 1991CAT 291	2	229			273					265	267	707 786	258	775	
240 240 314 314 200 201 201 201 285	297	229 234	235			265	251	266	293	200	216	271 271	273	200	
275 249 214 214 200 201 201 201 201 205 302	297 246 251	239 234 295	24 2	287			257	310	201	266	223	305 200	304	273	
249 249 314 117000 117000 201 201 285 302 296	297 246 251 276	229 234		287	717	264						274 247	284		
249 249 314 314 200 201 201 201 205 302	297 246 251	239 234 295	24 2			764								200	
249 249 214 214 200 201 201 201 201 205 312 296	297 246 251 276	239 234 295	24 2			764								200	
249 249 314 17000 1991CAT 285 312 285 312 296	297 246 251 276	239 234 295	24 2			764						270 291	201	2=0	
275 249 314 ITCDOP 1997CAT 291 285 302 296 241	297 246 251 276 293	239 234 295 263	247	296		764						270 291	201	200	273
249 314 115000 116 201 285 302 286 241	297 246 251 276 293	234 234 245 263	247	296	7 [7							270 291 275 267 305 200 254	291 281 00FDA	200	273
249 314 11-000 1000 201 285 302 296 241	297 246 251 276 293	239 234 295 263	24 2	296 WOLFF STANDA	717 PD 051	VIATION						270 291 275 267 275 250 2017COND 254 1980[CATION] =	201 281 00504 7	280 797 797	273 DAVIES
249 314 11-000 1000 201 285 302 296 241	297 246 251 276 293 444 TICN =	234 234 285 263 263	243 797 797 4074	296 WOLFF STANDA 250	بان ۱۱۰ دی ۱۱۰	INTION 315	277	256	225	226	230	270 291 275 260 017000 254 1400[CATION = 257 275	201 281 00504 7 265	280 797 797 798 247	273 DAVIES STANDA 289
975 249 314 1175000 281 285 352 296 241 0010000 1000104 256 217	297 246 251 276 293 444 TICN = 273 273	239 234 295 263 00FD 9 306 273	243 797 4109 240 274	296 WOLFF STANDA 250 297	۲۱۲ ۲۱۹ ۲۱۹ ۲۱۹	714TION 315 290	277	256	302	342	2=0	270 291 775 247 705 254 705 254 705 275 775 275 273 225	291 281 00FPA 7 265 271	280 797 797 797 797 247 749	273 DAVIES STANCA 289 262
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	265	753	272	210	247	291	205	290	250	757	275	270	
	272	745	251	240	202	263	296	270	297	220	247	-=4	
		273										270	
	255		269	274	793	761	269	265	272	371	272		
	228	201	253	2=3	270	292	2=4	276	200	274	203	270	
	268	500	251	271									
		5.											
	AUTCORD		POFO	1100	DAVIES								
	TADICAT		1		STANCAS								
	797	2=3 .	275	767	757	200	308	274	501	747	229	794	
	277	2=0	279	294	273	272	292	254	247	251	295	254	
2	272	203	275	279		282	267	247	251	749	243	74 5	
	275	260	371	767	272	310	220	2=0	264	220	755	272	
	750	270	275	283	757	225	281	250	242	272	2=7	754	
	747	767	266	776	757	263	745	277	296	785	245	774	
	315	212	220	744	799	777	769	234	249	204	269	270	
	757	752	291.	793	271	255	263	240	275	317	292	273	
	784	259	240	202									
	ATTCOND	734	POFOI	TOP	WOLFF								
	IVADICAT	104 =	1		STANDAS	VEC OF							
	793	753	274	270	757	223	108	274	291	743	278	735	
	777	767	270	703	273	272	202	2=5	247	751	286	754	
	272	234	275	279	250	75.2	268	247	251	742	241	245	
	275	269	122	760	272	310	271	267	265	220	256	2.25	
	260	270	275	294	252	249	291	250	284	272	257	757	
1	243	267	266	225	257	2=2	313	315	274	24=	285	276	
	278	245	770	273		244	329	271	268	294	269	294	
	271	283	291	253	257	275	26C	264	255	314	282	270	
	294	240	238	203			-						
	CUTCOOD		OPED		DAVIES								
	THROTCAT		10		STANDA	עיבה חל	ATTOM	a DE	QEES				
	747	272	773	110	273	772	295	100	271	777	103	772	
	272	271	273	250	764	725	205	27!	300	107	293	723	
	307	200	289	104	273	271	241	233	277	276	288	290	
	297	225	257	275	235	226	253	270	275	274	264	274	
	277	277	200	272		271	270	270	287	207	200	717	
	304	200	373	257		724	208	274	273	277	205	774	
	274	767	284	784		7	744	203	274	760	272	767	
	201	201	270	794	272	224	797	205	200	202	291	730	
	776	769	290	292	(e) 1/(s)						00.00		
	TITCOOD		-PEO.		FRAT	C ~ 11							
	vabical		11	A 111.00	STANDA		ATTON	8 75	REFES				
	107	200	271	200	200	744	250	272	271	271	272	174	
	271	271	100	295	272	273	331	272	271	74 1	793	232	
	100	200	297	207	773	271	241	292	277	274	297	231	
	207	226	258	775	235	225	757	277	2 . 7	797	273	77-	
	271	271	273	200		274	273		276	>74	200	314	
		200	304	272		723	208	274	247	. 7 . 7	205	276	
	105	267	284	200	201	100	200	200	287	20%	772	794	
	274					270	277	292	264	2= 2	272	235	
	270	291	201	2=0		2:3	211	272	644	6			
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	TUTCOOP		ADED.	ATTR	DAVIES								
	INBOICA		7		STANCA							727	
	757	775	255	247		766	204	247	247	743	267		
	272	276	271	240		772	297	101	310	276	707	75.2	
	276	76.7	10	247	284	203	281	275	281	276	271	27-	
	284	274	277	300	329	319	297	297	275	294	272	100	
	258	25.0	257	267	245	245	262	270	311	271	275	752	
	273	275	267	270	240	272	107	201	276	273	2 * 3	285	
	271	27'.	250	272	769	273	294	200	267	756	173	313	
	377	727	262	755	306	272	269	295	274	100	312	274	
	791	274	200	775									

A IT COCO	175	ADED	ATOP	PAVIES								
THESTEAT		15		CTANDA	on nev	INTION	4 05	REES				
295	205	272	281	281	295	294	278	284	777	286	296	
297	798	777	294	207	170	300	275	302	100	200	770	
277	104	200	791	273	777	202	271	260	795	203	705	
205	760	273	202		322	280	299	284	277	244	303	
272	106		2-2	2 4.7	7 6	1-9	6-7		2.	244		
AUTOPOP		-050		PAVIES								
THROTCAT		14	41.4	STANDA			4 05	OFFS				
794	101	273	291	271	100	292	300	288	275	293	777	
200	797	275	260		203	293	241	236	291	276	212	
277	100	274	275	277	244	290	277	270	202	750	759	
204	200	294	244	171	200	75.8	294	291	231	253	772	
799	270	2	2.4		2	1-0	2.14	271	2 - 1	6 - 2		
AUTOORD		-010	TOP	PAVIES								
IVODICAT				STANCA		ATTON	16 75	DEEC				
203	200	297	757	202	749	221	275	221	269	302	776	
274	294	· · · · ·		774	230	257	290	287	770		274	
	244	261	213	114	240	257	241	247	114	277	2:4	
200												
CUTCOD		OFD.	A100	DAVIES								
TABLEAT		12		STANDA								
217	107	304	214		244	266	259	202	240	258	104	
271	506	208	501	304	309	308	303	289	29=	300	285	
794												
CUTCOCD		CDE 3	ated	DAVIES		-						
IMBDICAL		20		STANCA				GEES				
272	204	205	785	117	202	775	SOF	299	704	207	797	
287	300	298	200	207	243	725	301	304	200	290	303	
303	297	280	202	302	2 - 3	294	304	290	277	295	299	
293	799	2.02	205	203	297	297	294	290	321	201	277	
304	242											
CUTCOOP	224	CDED	ATCR	DAVIES								
THROTCAT		4		STANDA	יזה הס	POTTON	1: 00	OFFS				
314	751	765	778	797	222	797	294	796	766	201	797	
781	202	755	777	275	770	707	205	200	775	2 . C	770	
777	759	292	301	203	273	312	303	307	7.54	288	721	
374	273	286	275	272	292	267	317	272	247	264	255	
274	270	279	279	311	293	799	295	249	247	274	769	
273	764	251	240	107	760	260	292	272	742	273	754	
793	746	219	206	779	795	760	765	271	270	271	270	
777	790	252	760	279	202	273	200	276	2= 1	202	250	
787	209	750	771									
AUTCOAD	775	ADED!	TOD	FRAFOT								
TVERTCAT		7	10-10 U.	STANDA		INTION	10 05	PFFS				
257	305	282	797	292	766	297	297	298	724	289	279	
265	252	314	204	302	203	270	272	270	202	275	733	
296	294	270	206	279	774	112	304	379	255	288	722	
304	274	797	275	272	793	203	312	270	770	274	256	
265	263	272	210	247	291	205	290	250	767	275	270	
272	745	251	240	302	763	296	270	297	770	247	-=4	
265	273	269	774	793	761	269	265	277	371	272	270	
											270	
228	201	253	253	270	292	2=4	276	290	274	203	214	

AUTCORD 254	-2521	901		5.74							OUTCOOD	35.5	OPFO	TOP	FRAFRT	SON						
INDDICATION :	-		\$7.1	NAU ULA	11111	10 05	-0554				IMROICA	-	- 2		STANDAS	TO DEV	TATION	11 25	PFFS			
764 797	275	795		794	747	310	267	276	7=1	797	276	756	. 775	250	297	799	7 25	224	225	776	257	264
276 300	175	270	- = 4	274	777	308	372	317	7 . 7	297	771	775	767	747	275	766	744	249	767	200	270	273
276 204		100	203	267	744	274	254	257	2-5	271	705	764	254	797	269	787	771	223	252	767	295	775
270	262	265	271	266	257	258	250	244	294	247	295	201	255	240	208	284	223	265	2	268	273	226
749 249	26*	282		227	271	269	242	278		300	306	273	769	273	270	745	274	740	240	770	228	273
271 259	27:	271	295	273	273	276	201	307	272	239	286	274	279	263	296	252	240	264	275	230	766	265
249 266	741	272	743	273	223	299	240	297	303	313	240	749	739	252	256	745	243	296	252	200	300	210
17 701		204	2.66	245	244	112	241	214	2-1	114	280	202	250	226	240	773	755	248	307	310	317	. 203
011TC000 764			-								264	264	-050	141								
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379 767	23:	270	207	295		221	22-	244	2=5	277	245.	791	200	778	247	756	777	240	270	775	249	212
260 267	24.3	294	7 8 3	271	2 - 5	200	220	272	272	270	280	276	270	100	240	774	307	274	270	316	272	296
174 275	3 1	259	247	293	274	274	2	232	2+4	291	207	ACF	105	298	241	795	224	275	277	297	279	763
760 767	74 7	700	721	220		270	277	746	772	25 2	794	26 R	751	263	296	293	177	203	202	294	203	3.20
753 745		201	7 . 4	207		209		744	749	270	201	760	790	294	205	299	174	303	270	227	204	292
270 269	774	760	2.2.2	260	2	270	2= 2	776		794	266	296	280	220	276	280	200	201	277	200	310	280
010 120		200		775	- 24	272	2	272	243	764	101	717	277	233	310	296	200	262	200	775	295	278
764 202	275	764	752	234	205	310	260	249	277	267	202	777	744	776	777	258	294	257	288	500	274	250
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276 233	279	305	791	297	2.46	278	27=	276	277	302	242	253	243	262	260	274	273	274	200	285	272	257
	274	269	747	277	793	284	243	276	275	267	788	295	270	730	255	101	200	270	279	715	274	243
771 767	741	777	747	791	249	249	375	745	744	296	767	267	256	232	747	254	746	219	274	226	240	
313 370	725	294	795	269	767	264	777	767	112	262	755	777	272	254	272	299	222	767	742	767	2=4	779
789 795		317	2	267	100	200	727	777	> = =	267	294	760	267	269	272	777	260	311	294	755	253	235
222 270	747	200		797	236	200	747	200	2=1	245	253	267	242	272	747	239	290	261	250	255	259	264
750 202	2 - 2	269	274	275	276	291	289	795	2=1	250	230	744	245	222	284	243	2=4	272	295	254	245	284
707 705		777		6.11	4 14.		-			2.2.4	786	250	257	171	6.4.4	2.4.2	2 4	e . c			1.4.2	104
OUTCOOD 294	-07011		101 ==								AUTCOOD	785	-000		PAVIES							
TURDICATION -	-			n ney	11	12 05					Inpoles.	-			STANDAR	DA DEV	ATTON	12 75/	DEEC			
780 706	7 2 7	745		291	727	312	270	200	114	25 9	764	204	706	7=4	791	761	22-	275	277	305	200	251
747 271	2.1	777	747	267	772	267	224	220	2 7 3	774	285	771	212	744	319	291	211	278	300	750	290	261
705 706	272	726	767	297	777	294	24-	748	2 = =	770	254	710	275	745	775	740	294	774	221	200	794	312
250 275	27:	311	243	263	234	314	277	2 9 1	240	341	277	769	275	201	260	280	297	299	200	791	291	749
231 285	297	200	279	265	797	245	274	705	2-7	301	200	276	293	275	310	2=3	281	275	293	255	259	272
203 250	707	274	715	312	260	240	3-3	220	251	270	260	212	200	2=4	256	767	775	295	256	707	227	237
768 272	23=	270	277	278	233	304	253	280	2=3	254	269	217	777	214	200	767	200	225	201	300	209	290
778 797	793	200	747	200	200	300	271	783	307	319	268	225	289	212	212	310	261	316	277	759	248	277
707 744		707									201	202	202	274								
ADC 000-1110	-00041		1.==								AUTCOOD	201	PED.		ECCEPTS							
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704 767	243	774	774	291	795	30.4	2-1	240	203	291	287	252	204	305	763	240	279	223	300	305	289	260
221 223	2 - 2	777	794	744	2.3	2 = 1	272	229	243	270	283	269	310	243	723	290	210	276	270	259	291	274
700 251	271	288	772	255	376	277	275	757	779	277	257	275	310	282	796	227	223	284	247	275	263	273
799 300	729	291	235	271	701	304	724	746	7= 4	300	299	252	260	273	307	272	279	773	295	253	219	247
777 777	202	700	74-	777	277	770	771	761	704	260	250	312	708	267	755	209	701	777	207	766	257	312
767 771	764	779		757	794	254	2 - 7	100	123	256	275	315	266	227	236	295	254	294	274	745	207	288
762 272	742	251		269	2=3	264	2-=	271	175	294	266	272	789	110	710	100	114	267	275	256	744	775
246 252	262	271		244	. ,	2.1.4	¢.				780	754	201	272				e		4. · · ·		
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The Constant of the Constant	-2523		-		ATTON	13 CF							0050 5	\$01 <i>0</i>	STANDAS	יאפר רי	INTION	A DEC	REFS	• •		
INDDICATION #	7			In nEV	272	13 DF 345		250	290	298				299			277	A DE7	274	290	2	285
209 315			STANDAS					250	207	285	IMBOICAT	-	5		STANDAS	279 256	277	295	274	271	272	246
209 315	7	280	5745045 271	248	272	345	341	205 741	297 250	285 319	10001071	273	5 277 263 297	299 245 286	STANDAS 290 220 251	279 279 256 271	277 271 271	295 295 267	274 287 272	271 257	272 275	246
10001CATION = 299 315 290 277	242	280 296	271 272	248 274	272	345 289	341 277 254 274	205 341 280	297 250 745	285 319 269	14001CA1 203 272	273 273 273 273 262	5 277 263 297 305	200 245 286 203	STANDAS 290 220 251 292	279 279 256 271 291	277 271 271 242	295 296 267 273	274 287 272 275	271 257 274	272 275 272	246 241 296
299 315 299 217 290 277 244 260	7 242 271 291	280 296 289	271 272 273	248 274 300	272 294 301	345 289 276	341 277 244 274 309	205 341 280 272	297 250 745 219	285 319 269 261	1980 (CA 203 272 276 281 274	273 273 273 262 271	5 277 263 297 305 272	299 245 286 293 277	STANDAS 290 220 251 292 273	279 276 276 271 291 300	277 271 271 248 248	295 295 267 273 267	274 287 272 275 788	271 257 274 296	272 275 272 287	246 241 296 311
14001CATION = 249 315 744 260 314 301	7 242 271 291 291	280 296 299 279	271 272 273 273	248 274 300 261 277 250	272 294 301 279 294 279	345 289 276 292 272 294	341 277 254 274 309 275	295 341 280 272 272	297 250 745 219 274	285 319 269 261 310	1000 (CA 203 272 276 281 274 269	273 273 273 273 262 271 269	5 277 263 297 305 272 247	299 245 286 293 277 274	STANDAS 290 220 251 292 273 310	279 276 271 291 289	277 271 271 248 301 277	295 295 267 273 267 273	274 287 272 275 275 288 270	271 257 274 296	272 275 272 287 270	246 241 296 311 268
10001CATION = 209 315 700 277 744 260 314 301 294 211	7 242 271 291 291 294 294 294 299	280 286 289 279 279 279 201 287 270	271 272 203 201 273 201 273 273 273	248 274 300 261 277 250 301	272 294 301 279 794 794 749	345 289 276 202 272 204 267	341 277 254 274 309 276 276	205 341 280 773 277 272	297 250 745 219 204 279	285 319 269 261 310 267	1000 (CA) 203 272 276 281 274 260 266	273 273 273 262 271 260 273	5 277 263 297 305 272 247 266	200 245 286 203 277 274 265	STANDAS 290 220 251 292 273 310 304	279 279 256 271 291 291 289 289 289	277 271 271 248 201 277 276	295 295 267 273 267 277 277 289	274 287 272 278 278 279 279	271 257 274 296 300 280	272 275 272 287 270 200	246 241 296 311 26P 270
209 315 700 277 744 260 314 301 274 211 259 276	7 242 271 201 202 202 203	280 296 299 279 279 271 297 270 270	271 272 203 201 273 201 273	248 274 300 261 277 250	272 294 301 279 294 279	345 289 276 292 272 294	341 277 254 274 309 275	295 341 280 272 272	297 250 745 219 274	285 319 269 261 310	1000 1041 203 272 276 281 274 266 266 291	273 273 273 262 271 260 271 260 273 251	5 277 263 297 305 277 247 266 294	200 245 286 203 277 274 265 240	STANDAS 290 220 251 292 273 310	279 276 271 291 289	277 271 271 248 301 277	295 295 267 273 267 273	274 287 272 275 275 288 270	271 257 274 296	272 275 272 287 270	246 241 296 311 268
209 315 740 277 744 260 114 301 764 211 764 211 764 214 744 217 744 214 752 246 237 312	3 242 271 241 241 244 244 244 244 244 244 244 24	280 296 299 279 279 271 297 270 270 269	STANDAG 271 272 273 273 273 273 273 273 277	248 274 300 261 277 250 301	272 294 301 279 794 794 749	345 289 276 202 272 204 267	341 277 254 274 309 276 276	205 341 280 773 277 272	297 250 745 219 204 279	285 319 269 261 310 267	1990 (CA 203 272 276 281 274 266 266 266 291 274	273 273 273 262 271 260 273 251 251 271	5 277 263 297 305 277 247 266 294 288	299 245 286 293 277 274 265 249 277	STANDAS 290 220 251 292 273 310 304 272	279 279 256 271 291 291 289 289 289	277 271 271 248 201 277 276	295 295 267 273 267 277 277 289	274 287 272 278 278 279 279	271 257 274 296 300 280	272 275 272 287 270 200	246 241 296 311 26P 270
209 215 209 215 200 277 244 260 314 301 260 276 200 276 200 276 252 246 237 312 000 000 314	3 242 271 201 303 204 203 203 277 205 277	280 296 299 279 301 287 270 270 270 269	STANDAG 271 272 273 273 273 273 277 277	248 274 300 261 277 250 301 279	272 294 301 279 294 253	345 289 276 292 272 204 267 297	341 272 244 274 274 274 274 274 275	205 341 280 773 277 272	297 250 745 219 204 279	285 319 269 261 310 267	1990 (CA 293 275 281 274 269 266 291 274 274	273 273 273 262 271 262 271 269 273 251 251 271	5 277 263 297 305 277 247 266 294	299 245 286 293 277 274 265 249 277	STANDAS 290 220 251 292 273 310 304 272 WOLEE	279 279 276 271 291 300 289 289 289 289	277 271 271 248 301 277 276 285	295 296 267 273 767 273 277 299 283	274 287 272 275 275 275 277 277	271 257 274 296 300 280	272 275 272 287 270 200	246 241 296 311 26P 270
200 215 200 215 200 217 244 260 114 301 274 211 274 27 252 246 237 312 001000 314 100000 314	7 242 271 701 700 740 700 207 207 207 207 207	280 296 299 279 279 270 270 270 269	STANDAS 271 272 273 273 273 273 277 277 277 277	248 274 300 261 277 301 279	272 294 301 279 294 269 269 269 253	345 289 276 292 272 204 267 297	341 272 244 274 274 274 274 274 274 275	205 341 280 273 272 272 280	207 250 245 219 294 279 291	285 319 269 261 310 267 266	1940 (CA) 203 272 276 281 274 266 291 274 274 274 01Corp 1940 (CA)	273 273 273 262 271 262 271 269 273 251 251 271 404	5 277 263 297 305 277 247 266 294 288 0052	200 245 286 203 277 274 265 240 277	STANDAS 290 220 251 292 273 310 304 272 WOLFE STANDAS	279 279 256 271 291 291 289 289 289 289 289 289	277 271 271 242 301 277 276 285	295 295 267 273 267 277 299 283	274 287 272 275 275 275 277 277	271 257 274 296 300 280 280 261	272 275 272 287 270 200 265	246 241 296 311 268 270 259
299 315 790 277 744 260 114 301 744 211 744 211 744 211 752 245 237 312 917000 314 1907(2110) 3	242 271 271 271 271 271 271 271 271 277 277	280 296 279 279 270 270 270 270 269 226	ST AND AS 271 272 273 273 273 273 277 277 277 277 277	248 274 300 261 277 301 279 301 279 207	272 294 301 279 264 263 263 263 263	345 289 276 292 272 204 267 297	341 272 264 274 276 276 277 277 277	205 341 280 273 272 272 280	297 250 245 219 204 279 291 291	285 319 269 261 310 267 266 215	1990 (CA) 203 272 276 281 274 266 266 291 274 0UTCOCO 1990 (CA) 274 254	273 273 273 262 271 260 271 260 273 251 271 404 10N = 253	5 277 263 297 305 277 266 294 288 0050 1 259	299 245 286 293 277 274 265 249 277 4709 254	STANDAS 290 251 292 273 310 304 272 WOLEE STANDAS 276	279 276 271 291 300 289 294 290 289 294 290	277 271 271 248 201 277 276 285	295 296 267 273 767 273 277 299 283	274 287 272 275 288 270 277 277 277	271 257 274 296 300 280	272 275 272 287 270 200 265	246 241 296 311 268 270 250
298 315 298 217 298 277 244 260 214 201 204 211 204 211 209 276 217 212 217000 314 209 2161104 = 248 269 247 291	3 242 271 701 700 740 700 700 707 70 707 70 707 70 707 70 70	280 296 279 279 270 270 270 270 260 226 283	ST ANDAS 271 272 273 273 273 273 273 277 277 ST ANDAS 275 275	248 274 300 261 277 360 301 279 207 263	272 294 301 279 294 269 269 269 269 269 269 271 271	345 289 276 292 277 297 264 267 297	341 272 264 274 276 276 277 277 277 277	205 341 280 273 277 277 280 306 269	207 250 245 219 294 279 291	285 319 269 261 310 267 266	1940 (CA 203 272 276 281 274 266 291 274 274 274 274 274 274 274 274	273 273 273 262 271 262 271 269 273 251 251 271 404	5 277 263 297 305 277 247 266 294 288 0052	200 245 286 203 277 274 265 240 277	STANDAS 290 220 251 292 273 310 304 272 WOLFE STANDAS	279 279 256 271 291 291 289 289 289 289 289 289	277 271 271 242 301 277 276 285	295 295 267 273 267 277 299 293 12 DFC 257	274 287 272 275 275 275 277 277	271 257 274 206 200 280 261	272 275 272 287 270 200 265	246 241 296 311 268 270 259
1000/74/104 209 315 200 277 244 260 214 260 274 211 274 211 274 211 274 214 237 312 245 245 237 312 247 293 247 293 248 249 247 293 248 249 248 249 249 249	242 271 271 271 271 271 271 271 271 277 277	280 296 299 279 279 270 270 270 270 270 270 270 270 270 270	ST AND AS 271 272 273 273 273 273 277 277 277 277 277	248 274 400 261 277 260 301 279 207 263 290	272 294 301 279 264 263 263 263 263	345 289 276 292 272 204 267 297	341 272 264 274 276 276 277 277 277	205 341 280 273 272 272 280	207 250 245 219 276 270 291 295 291 295 291	285 319 261 310 267 266 257 257 270	1990 (CA) 203 272 276 281 274 266 266 291 274 0UTCOCO 1990 (CA) 274 254	273 273 273 262 273 262 273 267 273 251 271 404 404 10N = 253 251 233	5 277 263 297 305 277 247 266 294 288 0050 1 259 283	299 245 286 293 277 274 265 249 277 277 277 277 277 277 277 277 277 27	STANDAS 290 251 292 273 310 304 272 WOLEE STANDAS 276 263	279 276 271 291 291 291 291 291 291 289 294 290 0 DEV' 277 275	277 271 271 248 201 277 276 285 285 (ATION 254 262	295 295 267 273 760 277 299 283 12 DFr 267 290	274 287 272 275 288 270 277 277 277 277	271 257 274 296 300 290 261 261	272 275 272 287 270 200 265 256 274	246 241 296 311 26P 270 250
298 315 298 315 290 277 244 260 314 301 274 211 274 211 274 217 274 274 237 312 237 312 247 245 247 247 247 247 247 247 247 247 247 247 247 247	743 743 744 764 765 777 755 755 755 755	280 296 299 279 279 270 270 270 270 270 270 270 270 270 270	CTANDAS 271 272 273 271 273 273 277 277 277 277 277 277 277 277	248 274 300 261 277 250 301 279 207 263 297 263 297	272 294 301 279 294 295 295 295 295	345 289 276 202 272 204 247 297 247 297 14 DF 300 265 250 273	341 272 244 274 274 274 274 277 247 271 273 247	205 341 280 273 277 280 306 269 280 306	297 250 245 219 270 291 291 291 295 320	285 319 261 261 267 266 215 257 270 239	10001(21) 203 272 276 281 274 266 266 266 274 017000 14001(21) 254 254 254 254 254	273 273 273 262 262 267 260 273 260 273 251 204 100 = 253 251	5 277 263 297 305 297 247 266 294 288 005 29 283 259 283 240	299 245 286 293 277 274 265 249 277 274 277 274 277 274 277 254 279	STANDAS 290 251 292 273 310 304 272 STANDAS 276 276 271	279 279 256 271 291 291 289 289 289 289 289 277 275 299	277 271 271 242 301 277 276 285 285 (ATION 254 262 243	295 296 267 273 767 277 289 283 12 DFr 287 290	274 287 272 275 275 277 277 277 277 275	271 257 274 296 300 280 261 273 261	272 275 272 287 270 200 265 256 274 272	246 241 296 211 268 270 250 250
1000/74/104 209 315 200 277 244 260 314 301 244 260 274 211 240 274 237 312 247 312 247 244 247 244 247 244 247 247 247 247	242 271 201 204 204 204 204 204 204 204 204 204 204	280 296 299 279 279 270 270 270 270 270 270 270 270 270 270	STANDAS 271 272 273 273 273 273 273 277 277 277 277	248 274 300 261 277 301 279 301 279 207 263 290 771 282	272 294 301 279 249 249 253 253 271 271 271 271 271 255 227	345 289 276 277 204 267 297 14 DE 300 265 265 265 273 259	341 272 264 274 276 276 277 277 277 277 277 277 277 277	205 341 280 272 272 280 306 269 280	2#7 250 745 219 279 291 241 285 273 241	285 319 261 310 267 266 257 257 270	1900 (14) 272 276 281 274 266 291 274 1900 (14) 274 1900 (14) 274 288 288 288 298	273 273 273 273 273 273 273 271 240 271 271 271 271 271 271 271 273 273 273 273 273 273 273 273 273 273	5 277 263 297 305 277 247 266 294 294 294 294 295 240 240 245	299 245 286 293 277 274 265 249 277 277 277 277 277 277 277 279 277 279 279	STANDAS 290 220 251 292 273 310 304 277 STANDAS 276 263 271 293	20 DEV 279 256 271 291 300 289 294 290 277 275 299 277 275 299	277 271 271 242 301 277 276 285 (ATION 264 262 248 273 272 270	295 296 267 273 277 277 277 277 287 283 12 DFr 287 287 287 287	274 287 272 275 277 277 277 277 277 277 277 27	271 257 274 286 299 261 261 277 261	272 275 272 287 270 260 285 274 274 274 274 272 287 287	246 241 296 11 26P 270 259 270 259
298 315 298 315 290 277 244 260 314 301 274 211 274 211 274 217 274 274 237 312 237 312 247 245 247 247 247 247 247 247 247 247 247 247 247 247	3 242 271 271 271 271 271 271 273 274 277 277 277 275 275 275 275 275	280 296 299 279 270 270 270 270 270 270 270 270 270 270	ST ANDAS 271 272 273 273 273 277 277 277 277	248 274 300 261 277 250 301 279 207 263 297 263 297	272 294 301 279 294 295 295 295 295	345 289 276 202 272 204 247 297 247 297 14 DF 300 265 250 273	341 277 264 274 375 275 267 275 277 277	205 341 280 272 272 280 306 280 306 280 280 280 280 271	207 250 245 219 274 274 274 274 291 285 320 273 241 247	285 319 261 310 267 266 215 257 270 239 264	1000 (CA) 2073 276 281 274 266 261 274 01(Coro 1400 (CA) 264 264 264 264 264 277	273 273 273 273 273 273 271 240 271 240 271 251 251 251 251 251 251 251 251 251 25	5 277 263 297 305 277 247 266 294 294 294 295 297 240 240 246 240	299 245 286 293 277 274 265 247 277 270 254 254 279 254 279 254 279 254 279 254	STANDAG 290 250 251 292 773 310 304 277 WOLEE 5TANDAG 276 263 271 293 293 293	20 DEV 279 256 271 291 291 291 299 289 289 289 277 275 299 269 269 269	277 271 271 242 301 277 276 285 (ATION 264 262 248 273 270 271	295 296 267 273 267 277 283 283 283 283 283 287 287 287 287 287 287 287 287 287 287	274 287 275 275 276 277 277 277 277 277 277 277 274 240 238 271 286 240 238 271 264 261 250	271 257 274 300 280 241 241 241 241 243 243 243 243	272 275 272 287 270 260 285 274 274 274 274 272 276 274 276 274	246 241 26P 270 250 270 250 270 250 270 254 274 294 274
209 315 209 315 200 277 244 260 314 301 274 211 274 211 274 211 274 214 237 312 246 273 347 293 247 293 247 293 249 275 247 293 249 277 249 277 249 275 249 277 249 274 249 277 249 276 249 277 249 276 249 276 249 277 249 276 249 276 240 277 240 276 240 277 240 276 240 277 240 276 240 277 240 277 240 276 240 277 240 27	3 242 241 241 241 243 244 244 244 244 244 244 244	280 296 299 279 270 270 270 270 270 270 270 270 270 270	CTANDAS 271 272 273 271 273 273 273 277 277 277 277 277 277 277	248 274 300 261 277 260 301 279 207 263 290 207 263 290 275	272 294 301 279 294 249 253 253 253 271 271 271 271 271 271 271 271 272	345 289 276 277 204 247 297 14 DF 300 265 265 265 259 259 259	341 277 274 274 275 277 277 277 277 277 277 277 277	205 341 280 277 277 280 306 280 306 280 280 280 280 280 280 280 280 280	2+7 250 245 270 270 270 271 241 245 320 273 241 247 277	285 319 261 310 261 310 267 266 215 257 270 239 254 239	1 vao (CA) 203 272 276 281 274 266 201 274 017000 1 vao (CA) 264 264 284 294 294 294 294 294 294 294 294 294 29	1104 = 273 273 273 262 271 267 271 267 271 271 271 271 271 271 273 271 273 273 273 273 273 273 273 273	5 277 263 735 277 247 247 294 294 294 294 294 294 294 294 294 294	299 245 286 297 277 274 265 249 277 270 277 254 270 299 299 275	STANDAG 240 250 251 242 273 310 272 WOLEE STANDAG 276 276 276 271 243 271 243 271 293 271 293	20 DEV 279 256 271 291 291 291 299 289 289 289 277 275 299 295	277 271 271 242 301 277 276 285 (ATION 264 262 248 273 272 270	295 296 267 273 277 283 277 283 283 283 283 284 287 287 287 287 287 287 287 287 287 287	274 287 275 275 276 277 277 277 277 277 277 277 277 277	271 257 274 260 280 241 251 273 261 273 263	272 275 272 287 270 260 285 274 274 274 274 272 287 287	246 241 296 111 270 259 270 259 270 259 270 259 270 259 270 259
1000/24/101 200 217 200 217 244 260 314 260 314 261 274 211 274 211 274 211 274 212 274 212 274 245 247 247 247 247	7 242 271 271 272 273 274 274 274 274 274 274 275 275 275 275 275 275 275 275	280 296 297 277 277 277 277 276 276 283 276 276 275 216 275 216 275 275 275 275	CTANDAS 271 272 701 270 277 277 277 277 277 277 277 277 277	20 DEV 248 274 300 261 277 260 301 279 207 263 279 207 263 271 282 275 275	272 294 301 279 794 749 253 749 253 749 253 749 253 271 271 271 271 271 271 271 271 271 272 726	345 289 276 272 204 267 297 14 DF 300 265 265 265 273 259 249 254	341 272 274 274 275 277 277 277 277 277 277 257 257 271 273	206 341 290 272 272 277 277 280 306 280 306 280 325 271 236 271 236	2+7 250 745 204 270 241 245 320 273 241 247 277 279	285 319 261 310 261 310 267 266 215 257 270 239 264 239 264 239 297	(va o (c a) 203 275 276 281 274 246 246 246 247 274 247 247 247 267 207	1104 = 273 273 262 261 262 271 262 271 267 273 273 273 273 273 273 273 27	5 277 2637 305 2777 266 294 294 1 259 240 249 249 249 249 249 249 249 249 249 249	299 245 286 293 277 274 277 274 277 277 277 277 254 277 290 275 280 275 280 244	574ND45 290 251 292 373 310 304 277 304 276 263 271 293 271 293 271 293 317	20 DEV 279 256 271 291 291 292 294 293 294 293 275 299 265 264 224	277 271 271 242 301 277 276 285 (ATION 264 262 248 273 270 271	295 296 267 273 267 277 283 283 283 283 283 287 287 287 287 287 287 287 287 287 287	274 287 275 275 276 277 277 277 277 277 277 277 274 240 238 271 286 240 238 271 264 261 250	271 257 274 300 280 241 241 241 241 243 243 243 243	272 275 272 287 270 260 285 274 274 274 274 272 276 274 276 274	246 241 26P 270 250 270 250 270 250 270 254 274 294 274
1000/74/101 209 315 200 277 244 260 214 260 214 201 204 211 204 211 204 211 204 214 217 214 217 214 217 214 247 217 247 217	5 242 271 271 271 271 271 274 277 277 277 275 275 275	280 294 274 277 270 270 270 270 270 270 270 276 276 276 275 275 275 275 275 275 275 275 275 275	CTANNAS 271 273 201 273 201 273 273 273 277 277 277 277 277 277 277	20 0EV 248 274 300 261 277 260 301 279 207 267 267 267 271 282 276 277 270	272 294 301 279 724 259 259 271 271 271 271 271 275 227 272 275 226 296	345 289 272 272 272 272 272 272 272 272 273 297 265 273 259 249 259 259 259 259	341 2754 2764 2764 276 276 277 277 277 277 277 277 277 277	206 341 290 272 272 277 277 280 306 280 306 280 325 271 236 271 236	2+7 250 745 204 270 241 245 320 273 241 247 277 279	285 319 261 310 261 310 267 266 215 257 270 239 264 239 264 239 297	1 vao (CA) 2073 276 2781 274 276 749 276 749 276 741 274 276 264 284 287 297 297 297 297 297 297 297 29	104 = 273 273 262 273 262 273 267 273 273 267 273 273 273 273 273 273 273 27	5 277 263 737 75 277 266 294 294 294 294 294 240 240 240 240 240 240 240 240 240 24	299 245 286 293 277 276 277 276 277 276 277 254 277 254 277 254 277 254 277 275 269 275 269 275 260 270	STANDAS 240 220 251 242 273 310 304 272 273 274 274 274 273 317 5565575	20 DEV 279 256 271 291 291 291 294 290 277 275 294 290 277 275 294 290 277 275 294 295 294 295 295 295 295 295 295 295 295	277 271 242 242 242 276 285 247 254 264 264 264 277 270 271 299	295 296 267 273 760 273 280 283 283 285 285	274 287 275 275 278 277 277 277 277 277 277 277 277 277	271 257 274 300 280 241 241 241 241 243 243 243 243	272 275 272 287 270 260 285 274 274 274 274 272 276 274 276 274	246 241 26P 270 250 270 250 270 250 270 254 274 294 274
209 215 209 215 200 277 244 260 214 201 274 211 274 211 274 214 275 245 277 312 247 312 248 245 247 312 248 245 247 291 248 275 249 275 249 275 249 275 249 275 249 275 250 255 251 100 234	3 242 771 201 304 203 203 203 203 203 203 204 204 203 203 203 203 203 203 203 203 203 203	280 2949 279 201 270 270 270 270 270 270 270 281 235 216 283 274 235 216 274 274 274	CTANDAG 271 272 273 273 273 273 273 273 275 275 275 275 275 275 275 275 275 275	20 DEV 248 274 300 261 279 207 263 297 271 282 275 277 270 270 270	272 294 301 279 724 259 259 271 271 271 271 271 275 227 272 275 226 296	345 289 272 272 202 277 297 297 297 297 297 297 297 259 245 273 259 249 263 259	341 272 264 274 274 274 274 277 277 277 277 277 27	208 341 290 277 277 277 280 306 280 306 280 280 280 280 280 280 281 283 283 284	2+7 250 745 204 270 241 245 320 273 241 247 277 279	285 319 261 310 261 310 267 266 215 257 270 239 264 239 264 239 297	1 VBO (CA) 203 272 276 781 274 276 276 276 276 276 276 276 276	104 = 273 273 262 273 262 273 267 273 273 267 273 273 273 273 273 273 273 27	5 277 2637 305 2777 266 294 294 1 259 240 249 249 249 249 249 249 249 249 249 249	299 245 286 293 277 276 277 276 277 276 277 254 277 254 277 254 277 254 277 275 269 275 269 275 260 270	574ND45 290 251 292 373 310 304 277 304 276 263 271 293 271 293 271 293 317	20 DEV 279 256 271 291 291 291 294 290 277 275 294 290 277 275 294 290 277 275 294 295 294 295 295 295 295 295 295 295 295	277 271 242 242 242 276 285 247 254 264 264 264 277 270 271 299	295 296 267 273 760 273 280 283 283 285 285	274 287 275 275 278 277 277 277 277 277 277 277 277 277	271 257 274 200 280 281 273 261 273 261 273 261 273 273	272 275 2747 270 270 270 275 276 276 276 276 276 276 276 276 276 276	246 241 26P 270 250 270 250 270 250 270 254 274 294 274
1000/74/104 209 315 200 277 244 260 214 260 274 211 274 211 275 274 211 275 274 211 275 275 275 275 275 275 275 275 275 275	7 242 741 743 744 754 754 755 754 755 755 755 755 755	280 2949 279 270 270 270 270 270 270 270 275 276 275	ST ANDAS 271 271 273 273 273 273 273 273 275 275 275 275 275 275 275 275	20 0EV 248 274 100 241 277 101 279 207 263 271 282 276 277 270 276 277 270	272 294 301 279 724 253 253 253 271 271 271 272 225 286	345 289 272 272 272 277 297 297 297 297 297 247 297 245 245 245 245 245 245 245 245 245	341 274 274 274 274 274 274 277 277 247 277 257 257 257 247 247 247 247 247 247	208 341 272 277 277 280 306 280 306 280 275 275 275 275 275 274 293 244	297 255 219 204 274 291 245 320 271 247 277 247 277 240	295 319 269 261 100 267 266 257 276 255 256 257 254 254 239 250 250	(veo (c.4) 203 203 275 276 276 276 276 276 276 276 276	104 = 273 273 273 262 271 240 271 240 271 271 271 271 271 273 271 273 273 273 273 273 273 273 273	5 2773 2947 105 277 264 288 2747 264 288 2847 264 288 276 218 276 218 276 218 276 276 7 7 7 7 7	299 245 286 293 277 265 249 277 265 249 276 279 290 276 280 276 280 244 100 298	STANDAS 240 251 242 251 242 273 310 304 273 277 243 277 243 273 317 55560555 273 317 55560555 273 273 277 273 277 273 277 273 277	D DEV 279 256 271 291 201 289 284 291 200 DEV 277 275 290 DEV 277 275 293 244 293 245 244 244 245 244 246 244 243 244 203 D 204 224	277 271 271 271 271 277 274 277 276 276 276 276 276 277 276 277 276 271 298	295 296 267 273 240 273 240 283 280 283 200 283 200 200 200 200 200 200 200 200 200 20	274 287 275 275 278 277 277 277 277 277 277 277 277 277	271 257 274 206 200 261 271 261 271 203 261 273 260 279	272 275 277 270 265 265 276 276 276 276 276 276 276 276 276 276	246 241 296 270 259 270 259 270 259 270 254 274 276 322 276
1000/24/104 200 215 200 215 200 214 200 214 200 214 200 214 200 214 200 214 200 214 200 214 200 214 200 215 201 205 201 205	7 242 771 701 703 743 704 704 704 704 704 705 741 704 705 741 705 741 705 741 705 741 743 744 755 741 745 745 745 745 745 745 745 745	280 298 278 277 270 270 270 270 270 270 270 270 276 276 276 276 276 276 276 276	CTANDAS 271 272 273 273 273 273 273 273 273 273 273	20 0EV 248 274 200 261 279 207 263 279 207 263 277 270 277 270 280 277 270 280 248 274 277 270 280 277 270 280 271 282 277 270 277 270 277 270 277 277	272 294 301 279 249 253 (ATICN 271 271 271 271 271 272 225 246 (ATICN 271 272 246 277 272 246 276 276 277 279 279 279 279 279 279 279	345 286 272 272 272 272 272 272 273 265 273 259 243 289 243 289 243 289	341 279 264 276 276 277 277 277 277 277 250 271 299 250 271 299 263 263 263 272	206 341 240 272 272 280 306 280 306 280 280 280 280 281 234 244 270	297 250 219 204 274 274 273 241 245 3203 241 247 273 240 240	295 319 269 261 261 257 256 257 275 257 275 233 264 233 234 237 250 277 272	(vac(c)) 203 203 275 276 281 274 243 244 244 244 254 254 257 257 257 257 257 257 257 257	104 = 273 273 273 273 267 271 267 271 260 271 260 271 261 273 267 273 273 273 273 273 273 273 27	5 2773 2647 3055 2773 2477 2447 2447 2447 2447 2447 244 272 2446 272 2446 2747 2446 2747 772 18 2773 2477 2477 2477 2477 2477 2477 2477	299 245 286 293 277 274 265 277 276 277 270 277 270 277 270 277 270 277 280 278 288 260 244 1700 298 266 293 277 280 293 274 203 277 274 203 277 274 203 277 274 203 277 274 203 277 274 203 277 274 203 277 274 277 276 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 274 277 277	STANDAS 240 220 251 242 273 310 304 272 273 274 274 274 273 317 5565575	20 DEV 279 274 271 291 291 292 294 293 295 277 275 296 244 293 244 293 244 293 245 245 254 224 305 254 224 305 275 277 275 275 275 275 275 27	277 271 242 242 242 276 285 247 254 264 264 264 277 270 271 299	295 296 267 273 760 273 280 283 283 285 285	274 287 287 276 277 277 277 277 277 277 277 277 27	271 257 274 200 280 281 273 261 273 261 273 261 273 273	272 275 2747 270 270 270 275 276 276 276 276 276 276 276 276 276 276	246 241 26P 270 250 270 250 270 250 270 254 274 294 274
1000/0017617671 200 315 200 277 244 260 314 301 274 211 274 211 274 211 274 211 274 211 274 211 1007/0016 277 312 274 245 274 245 274 245 274 245 275 327 266 275 201 207 275 327 276 275 201 207 276 275 276 275 277 26 276 275 276 275 277 26 276 275 277 26 276 275 276 275 276 275 277 26 276 275 276 275 276 275 276 275 276 275 277 26 277 26 277 26 278 275 276 275 276 275 277 26 277 26 278 275 276 275 277 26 277 26 277 26 278 275 277 26 275 276 275 276 275 277 26 275 276 275 277 26 275 276 275 276 275 277 26 278 275 276 275 276 275 276 275 276 275 276 275 276 276 276 276 277 276 277 276 277 277 277 276 277 277 276 277 276 277 276 277 276 277 276 277 276 277 276 277 276 276 277	3 242 241 241 241 241 241 244 244	280 296 297 297 297 297 270 270 270 270 270 270 270 270 270 27	CTANNAS 271 272 273 273 273 277 277 277 277 277 277	20 nev 248 274 300 261 277 260 301 277 263 290 276 276 277 270 276 276 276 276 276	272 294 301 279 249 253 749 253 749 253 749 253 271 271 271 271 272 224 286 141104 271 272 224 286	345 286 272 202 204 247 297 14 DF 300 265 247 273 259 249 249 249 249 249 249 249 249 249 24	341 274 274 274 274 274 274 274 274 271 271 271 271 241 241 241 241 241 241 241 241 241 24	206 341 277 277 277 280 306 280 306 280 326 280 326 271 234 244 276	207 255 216 2746 2740 2740 2740 2741 245 3747 2773 2467 2776 2764 2764	295 319 269 261 210 257 266 257 277 254 233 254 233 250 278 278 277 250	(1000 (CA) 2007 276 276 276 276 276 276 276 27	104 = 273 273 273 262 771 260 771 260 771 261 771 263 264 771 263 771 263 771 263 771 263 771 263 773 773 773 773 267 773 773 267 773 267 773 273 267 773 267 773 267 773 267 773 267 773 267 773 267 773 267 773 267 773 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 771 267 774 774 774 774 774 774 774 7	5 277 267 305 247 264 288 0050 1 259 244 264 276 218 276 218 276 218 276 218 276 218 276 7 787 7287	299 245 286 297 277 265 277 277 277 277 277 277 277 278 260 278 260 278 260 244 100 298 260 273	STANDAS 240 220 251 242 773 410 404 277 304 277 243 273 273 273 273 273 273 273 273 273 27	D DEV 279 256 271 291 291 201 289 284 200 DEV 277 256 260 264 205 264 264 264 264 264 264 264 264 264 274 275 264	277 271 271 271 271 277 276 276 276 276 276 276 276 276 271 298 (ATION 279 270 271 298	295 2967 273 267 273 287 287 287 287 287 287 287 287 287 310 287	274 297 297 275 275 277 277 277 277 277 277 277 27	271 257 274 207 274 277 274 277 274 277 276 277 276	272 275 272 277 270 200 200 200 200 200 200 200	246 241 296 111 250 270 250 270 301 274 274 274 274 274 274 274 274 274 275 206 273
100017417104 = 209 315 200 277 244 260 274 211 274 211 274 211 274 211 274 217 252 246 237 312 247 312 247 266 247 291 249 275 249 275 240 275	7 242 241 241 241 241 241 244 241 244 244	280 296 299 279 270 270 270 270 270 270 270 270 270 270	CTANNAG 271 272 273 273 273 273 277 277 277 277 277	20 nev 248 274 300 241 277 260 301 279 207 263 297 263 2971 282 277 275 277 270 280 243 276 243 276	272 294 301 279 749 253 749 253 749 253 749 253 271 255 227 272 226 296 271 259 271 259 271 259 271 259 272 272 272 272 272 272 272 272	345 289 270 272 202 267 297 14 DF 265 265 273 259 263 289 263 289 263 289 263 289 263 289 263 289 263 289 263 289 263 289 265 273 259 265 273 259 265 273 265 273 265 273 275 275 275 275 297 297 297 297 297 297 297 297 297 297	341 272 274 274 274 277 247 247 247 250 271 243 243 243 243 243	206 341 240 272 272 280 306 280 306 280 280 280 280 281 234 244 270	297 250 219 204 274 274 273 241 245 3203 241 247 273 240 240	295 319 269 261 267 265 257 279 239 266 239 259 259 259 259 259 259 259 259 259 25	1000 (CA) 203 272 276 281 274 264 274 274 274 274 274 274 277 277	104 = 273 273 273 273 267 771 260 771 260 771 260 771 261 771 261 771 261 771 263 771 264 764 764 764 773 267 773 267 773 267 773 267 773 267 773 267 773 267 773 267 773 267 771 267 772 275 775 275	5 2773 2747 3252 247 264 288 0 1 2583 246 246 278 2746 2747 246 2747 246 2747 7 721 265	299 245 286 297 274 265 277 276 277 254 277 254 277 288 264 277 288 264 279 276 288 264 279 276 288 264 279 276 276 277 276 277 276 277 277 277 276 277 277	STANDAS 240 220 251 242 773 310 773 304 777 244 273 271 243 271 243 271 243 273 273 273 273 273 273 275	D DEV 276 276 271 201 201 201 201 201 201 201 201 201 201 201 201 201 201 201 202 201 203 203 204 204 205 204 205 204 206 204 207 205 208 204	277 271 271 271 271 277 274 264 264 264 264 274 275 276 271 298	295 2967 273 267 273 267 273 287 287 287 287 287 285 11 055 285 11 055 285 285	274 287 287 276 277 277 277 277 277 277 277 277 27	271 257 274 100 241 241 241 241 241 241 241 241 240 240 240 240	272 275 275 270 270 275 270 275 275 275 275 275 275 275 275 275 275	2461 2461 266 270 250 270 250 270 250 270 261 280 276 322 280 278 322
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ORIENTATION MEASUREMENTS

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CUTCOD 24 OPERATOR WOLFE	PUTCOND 34 ODEDATOD EGGEDTSON	
VECTOR MEAN = 233 STANDARD DEVIATION 14 DEGREES	VERTOD VEAN = 235 STANDARD DEVIATION 22 DECRES	
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VECTOD VEA: = 237 STANDARD DEVIATION 16 DECREES		
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OUTGIND 46 DECATOR EGGETSON VECTOR VEAN = 21 STANDAD DEVIATION 16 DECRESS 228 244 215 214 219 240 242 241 197 211 215 217 243 202 214 240 214 215 217 243 210 214 240 214 240 214 240 214 212 210 256 207 233 217 215 216 212 214 216 241 157 149 269 227 257 268 185 215 241 157 149 269 227 257 266 185 216 241 157 149 269 227 257 266 185	239 235 235 244 257 214 240 223 245 227 193 227 233 255 257	NITROD 56 DESITO -115 VECTOR WEAN = 71 ST3+D2D DEVI+TION 10 DEGRES 716 728 101 721 739 740 725 273 234 739 750 715 730 727 730 715 748 730 734 734 746 740 713 105 720 726 721 727 710 724 723 742 720 734 744 744 744 744
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ANTCOND AS	POFDATAD	WOLFE								
VECTOR MEAN	793	STANDAS	Yar rev	TATION	TR DE	OFFEC				
273 234	224 244		236	216	194	271	104	194	215	
200 227	756 770		344	719	770	276	774	247	205	
272 264	750 184	251	311	743	200	275	710	278	741	
218 776	709 747		777	109	727	334	74=	247	219	
743 709	749 250	212	777	255	222	250	747	254	220	
745 77	116 750	716	795	202	276	197	222	202	197	
743 218	747 200		275	237	230	280	744	228	239	
244 235	203 232		217	235	219	239	200	252	187	
717 710	717 777									
	OPEDATOD	WOIFE								
VECTOD WEAN =		STANDAS	יזר הי	TATION	18 05	PEFS				
777 77?	759 237		750	711	249	253	21.0	355	214	
777 775	244 233		741	244	265	224	215	219	222	
250 220	216 274		213	211	274	277	795	226	181	
193 235	233 213		211	205	278	215	203	218	225	
101 203	192 190	243	243	237	242	358	107	207	228	
274 191	277 221	221	722	226	204	230	211	230	345	
250 243	250 255	249	2:0	218	272	179	221	223	239	
710 774	721 776	222	753	722	213	245	151	231	217	
763 773	744 775									
ALITCOCO 67	POPDATAD	MULEE								
VECTOR MEAN =	276	STANDAT	ישר רי	TATION	18 75	REFS				
228 215	244 230	229	238	249	240	244	149	222	192	
746 747	270 244	100	241	228	264	24=	214	203	230	
277 267	720 714	769	225	227		241	754	252	253	
197 271	777 750	200		232	24 -	203	220	256	230	
715 797	210 254		757	215	220	5= 2	777	219	216	
776 774	777 700	210	744	255	3=7	275	127	746	247	
213 26	744 773	725	222	231	211	247	777	255	202	
25A 242	257 190	101	775	278	217	240	1 = 1	210	295	
210 243	757 747									
OUTCOOD 69	POEPATOD									
VECTOD VEAN =		STANDAS								
204 204	735 277	241	245	273	222	244	4	257	196	
224 237	367 730	237	222	201	107	214	224	107	205	
250 204	774 775	741	220	226	228	100	203	210	238	
107 223	774 7:0		224	103	277	244	20%	206	256	
713 191	204 250		205	271	2.29	192	234	272	231	
231 244	237 196	255	255	236	241	236	247	246	254	
222 221	263 200		233	235	237	201	271	248	228	
240 255	235 100	275	235	288	278	261	273	230	191	
770 750	252 211									

